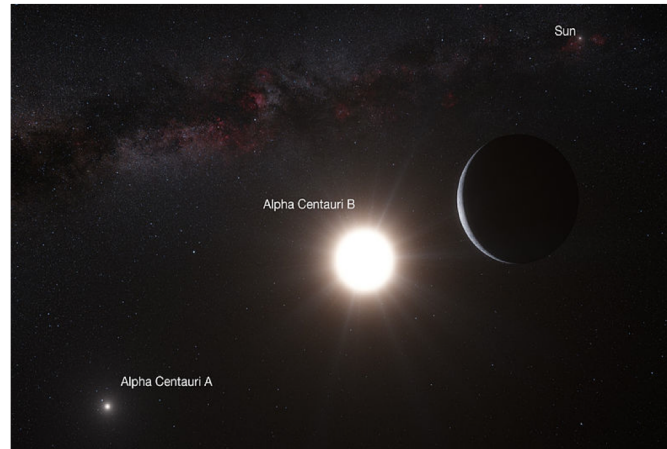
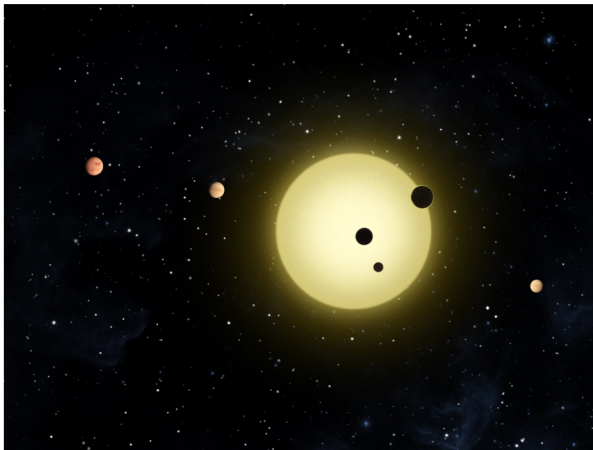


Introduction to exoplanetology

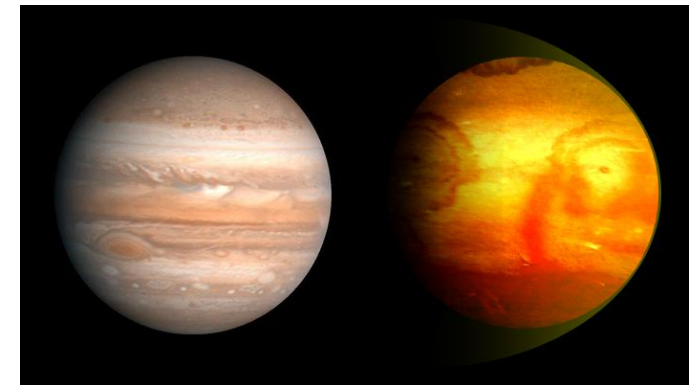
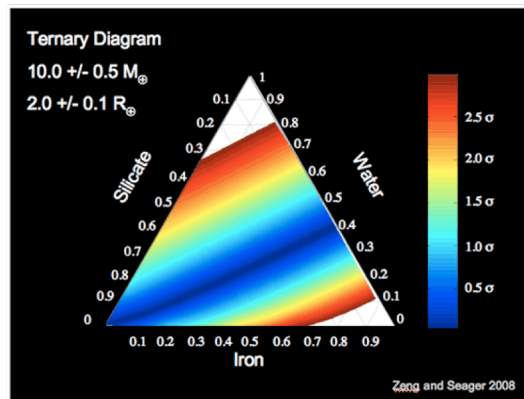
Michaël Gillon (michael.gillon@uliege.be)

Olivier Absil (olivier.absil@uliege.be)



Introduction to exoplanetology. IX.

Structure and atmosphere of exoplanets



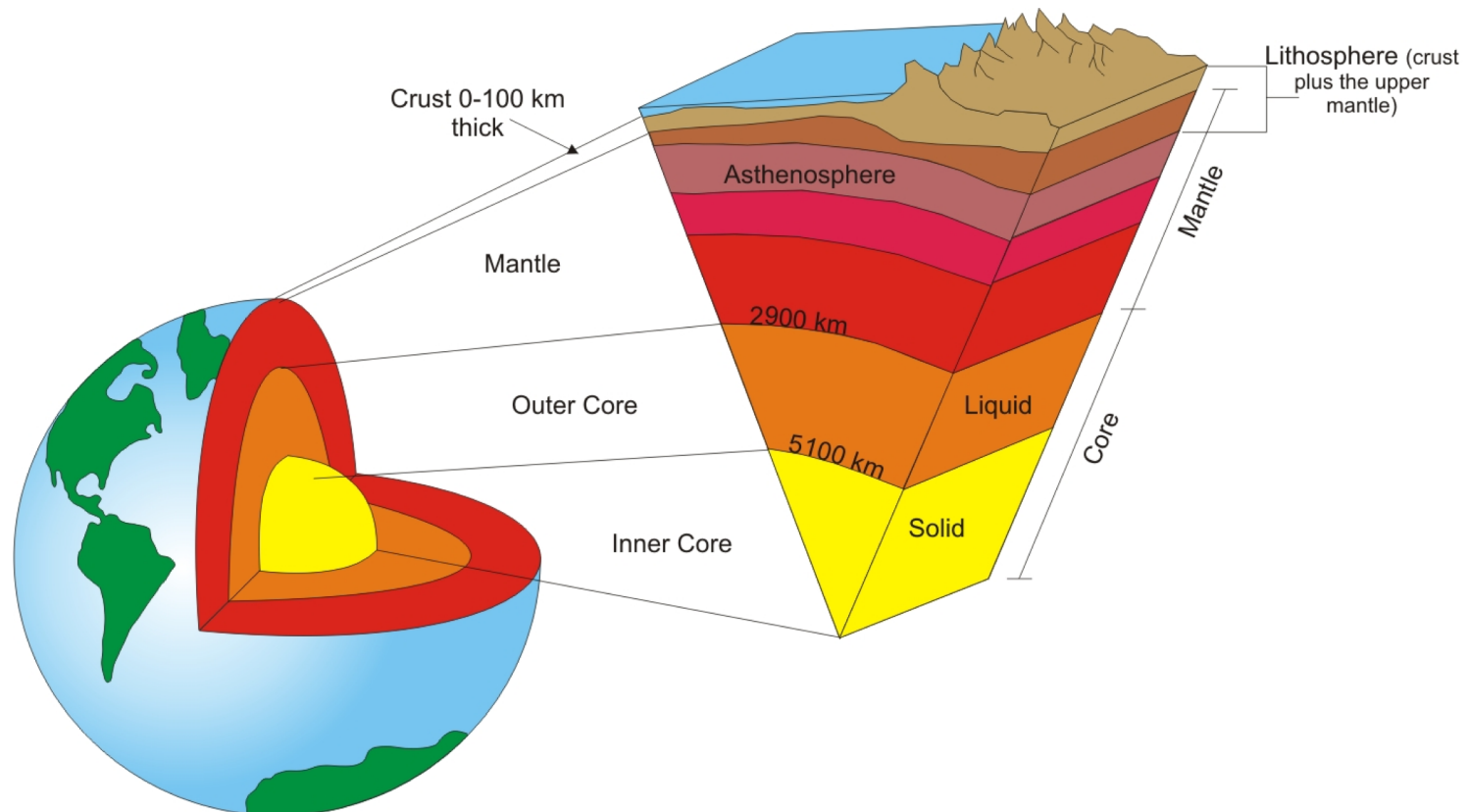
Michaël Gillon
michael.gillon@uliege.be

Internal structure of terrestrial planets

1. Earth

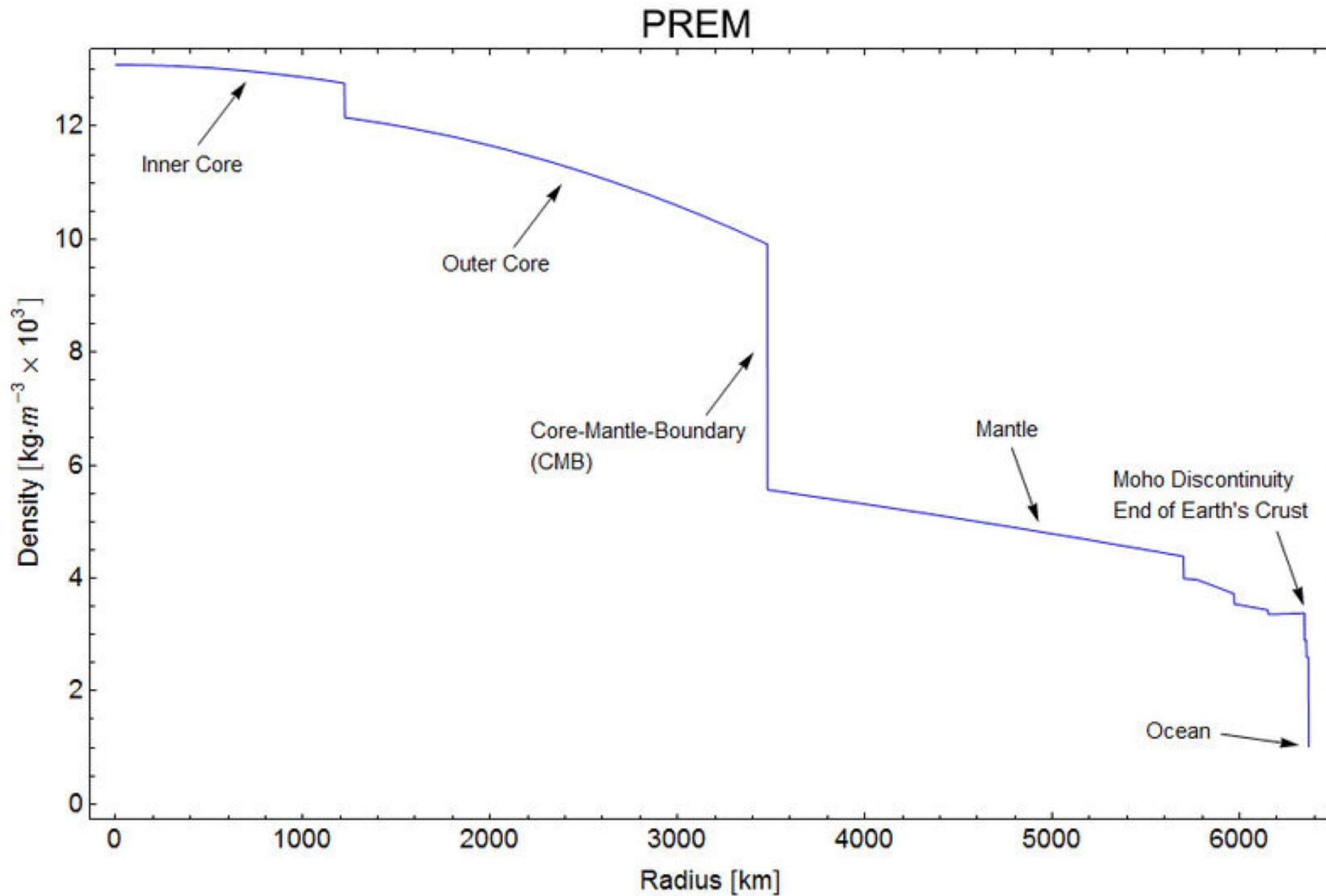
Differentiated structure:

Crust of silicates, viscous mantle of silicates enriched with Fe, outer + inner core of Fe



Internal structure of terrestrial planets

1. Earth

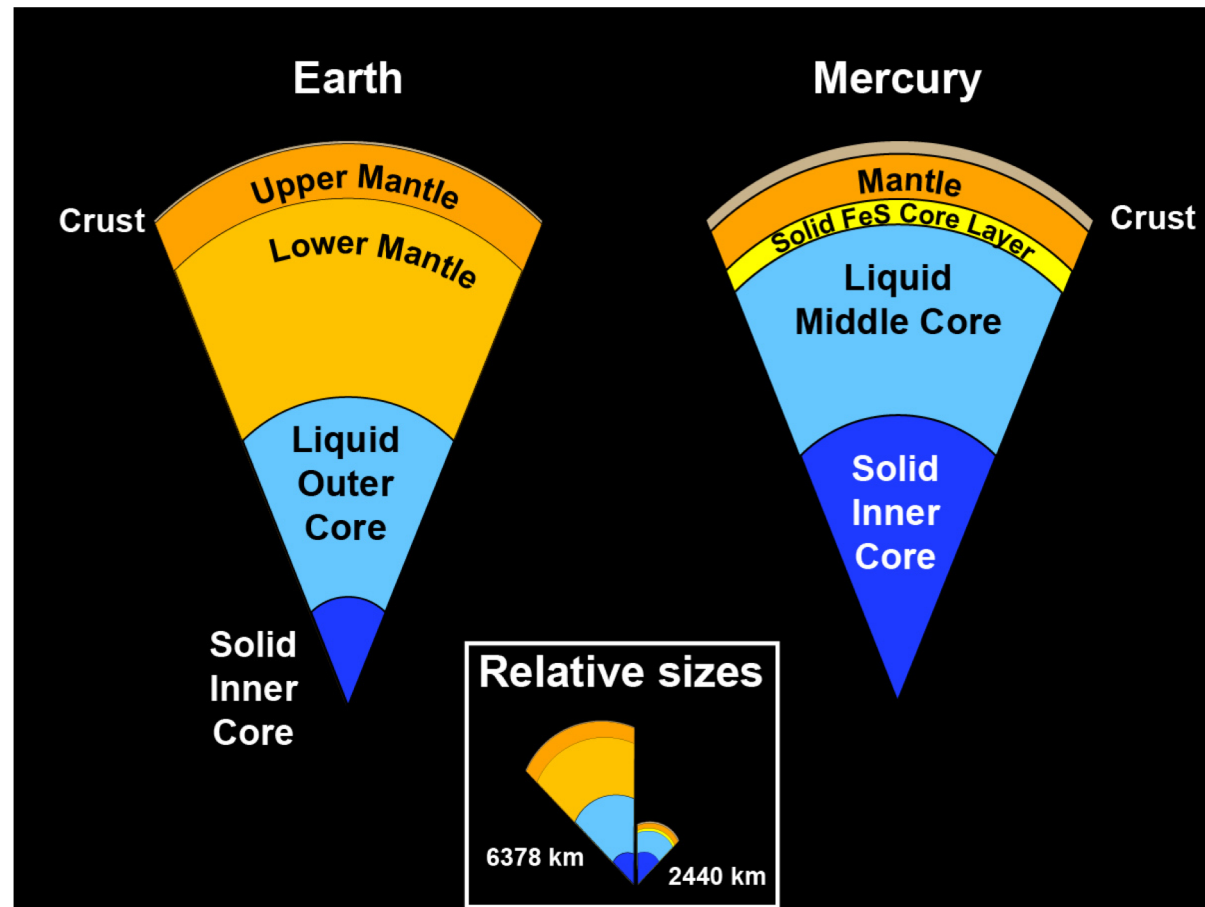


Internal structure of terrestrial planets

2. Mercury

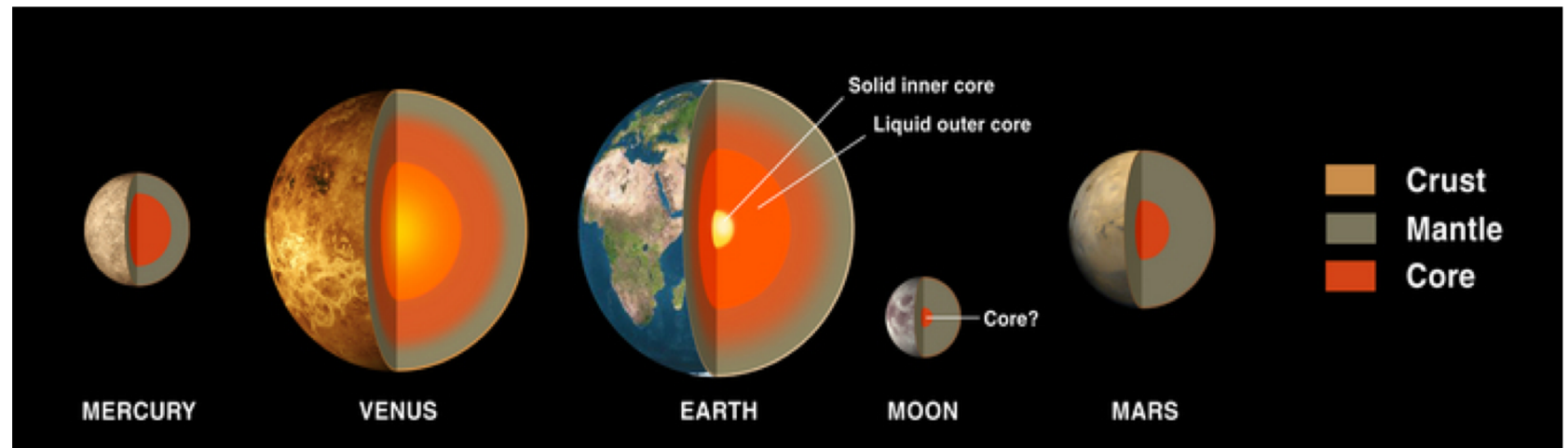
Significantly enriched in Fe compared to Earth

Origins? Maybe a giant impact having stripped out the planet of its outer shells
(Benz et al. 1988)



Internal structure of terrestrial planets

3. Mars, Venus, and the Moon



Differentiated structure: core – mantle - crust

| | | | | | |
|-----------------|-----|------|-----|----|----|
| Plate tectonics | no | no | yes | no | no |
| Volcanism | no | yes? | yes | no | no |
| Magnetosphere | yes | no | yes | no | no |

Internal structure of terrestrial planets

Elementary composition of the Earth

| | | | | |
|---------------------|---|--------------|---|--------------|
| O: 30.3% | | Ni: 2.0% | | |
| Fe: 33.4% | | Ca: 1.0% | | |
| Si: 19.2% | + | Al: 0.9% | + | rest... 0.1% |
| Mg: 12.2% | | S: 0.9% | | |
| Total: 95.1% | | Total = 4.8% | | |

Including other elements than **O, Fe, Si, Mg** changes the results of the models at a level <1%

Basic assumptions of terrestrial planet models

Planet composed of 2 basic components: Fe and silicates

Differentiated structure, with different elementary compositions for each layer

Earth: core, lower and upper mantle

crust? Negligeable. Same for *hydrosphere* and *atmosphere*

Internal structure of terrestrial planets

General structure equations of a planetary model

$$\frac{\partial P}{\partial r} = -\rho g$$

Hydrostatic equilibrium

$$\frac{\partial T}{\partial r} = \frac{\partial P}{\partial r} \frac{T}{P} \frac{d \ln T}{d \ln P}$$

Thermodynamic equilibrium

$$\frac{\partial M}{\partial r} = 4\pi r^2 \rho$$

Mass conservation

$$\frac{\partial L}{\partial r} = 4\pi r^2 \rho \left(\dot{\epsilon} - T \frac{\partial S}{\partial t} \right)$$

Energy conservation

$$\rho_i = f_i(P, T)$$

Equation of state (1 per layer)

+ boundary conditions

Internal structure of terrestrial planets

Modeling

Are assumed a planetary mass

O/Si, Fe/Si, Mg/Si, and Mg# = (Mg/Mg+Fe)_{silicates}

N layers

an elementary division for each layer (core: only Fe)

Elementary composition -> mineralogical composition

Structure equations

1. Mass conversation

$$M = 4\pi \int_0^R r^2 \rho(r) dr$$

with $\rho(r)$ depends on the local composition, T and P

Internal structure of terrestrial planets

2. Thermodynamic equilibrium (energy transfert)

Planetary interiors: transfer dominated by convection

$$\frac{\partial T}{\partial r} = \frac{\partial P}{\partial r} \frac{T}{P} \frac{d \ln T}{d \ln P} \quad \text{with} \quad \frac{dT}{dP} = \frac{\alpha T}{\rho C_p} \quad \text{For each layer, allowing for discontinuities at the junctions}$$

with α the thermal expansion coefficient (K^{-1}) and C_p is the thermal capacity (J/K)

3. Hydrostatic equilibrium

$$\frac{dP(r)}{dr} = -\rho(r)g(r)$$

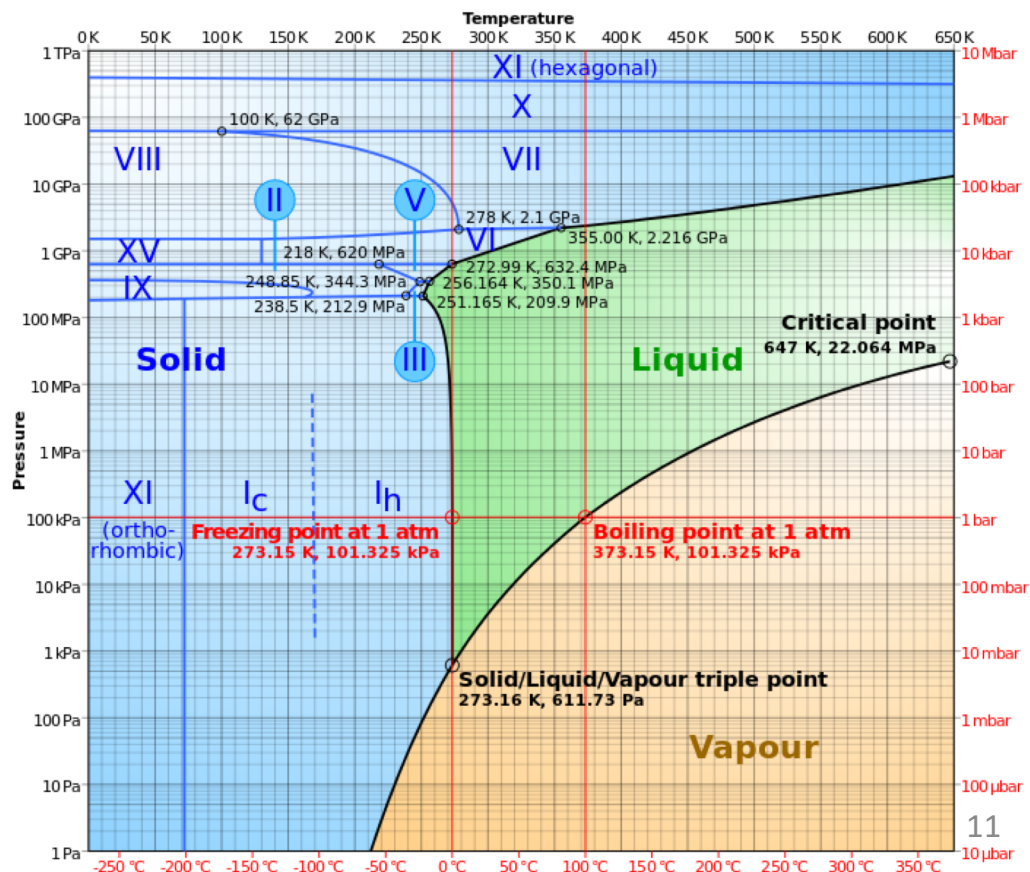
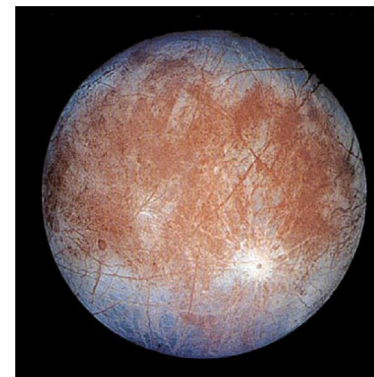
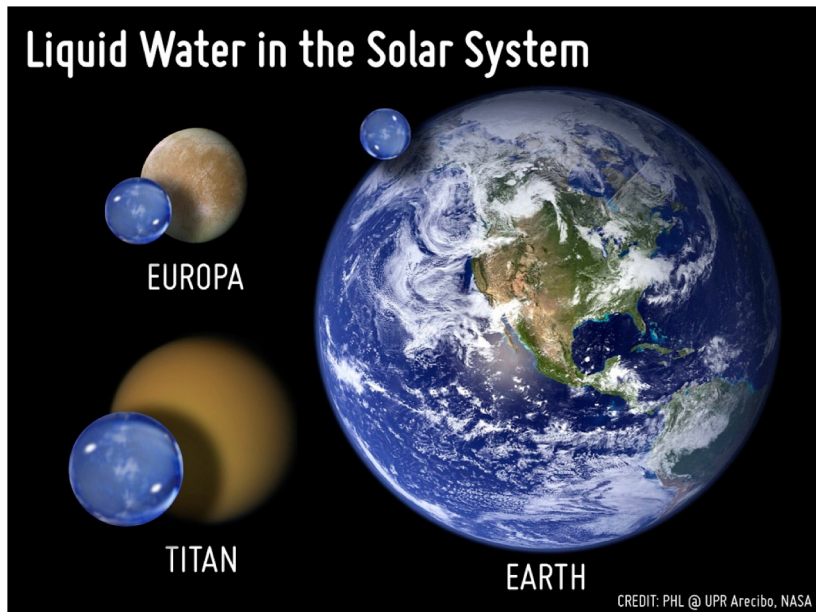
4. Equation of state (EOS)

$$\rho_i = f_i(P, T)$$

Well calibrated for Earth-mass planets, but not so much for super-Earths, for which pressures can be $> \text{TPa}$

Internal structure of terrestrial planets

Ocean planet: external layer of water



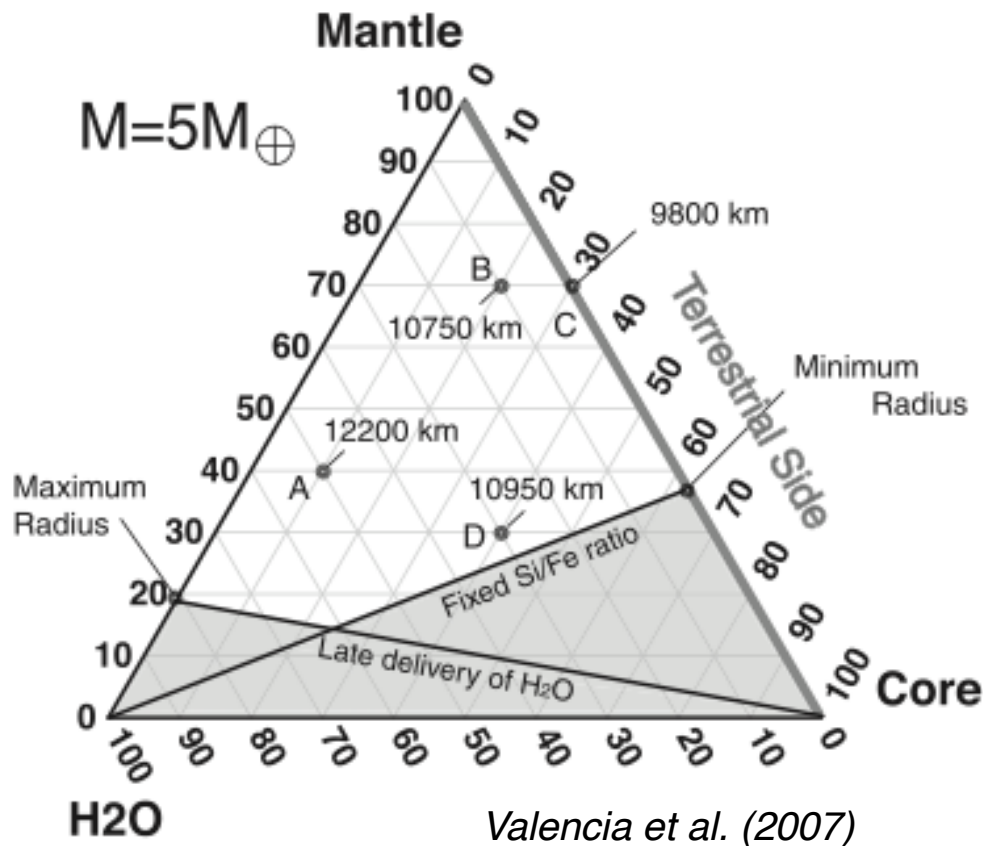
Water ice of type VII and X

Liquid water layer?

External thin layer of water ice of type I? (cold enough planet)

Internal structure of terrestrial planets

Ternary diagram

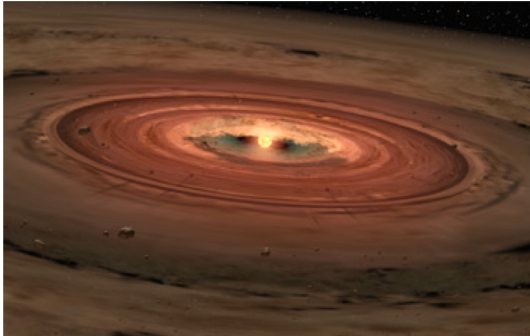


With M_p and R_p , the structure of the planet can't be unambiguously determined

Assume initial conditions: combining planetary and disk models?

Internal structure of terrestrial planets

Good idea a priori, but....



Planets can migrate by planet-disk, planet-planet, and planet/star interactions



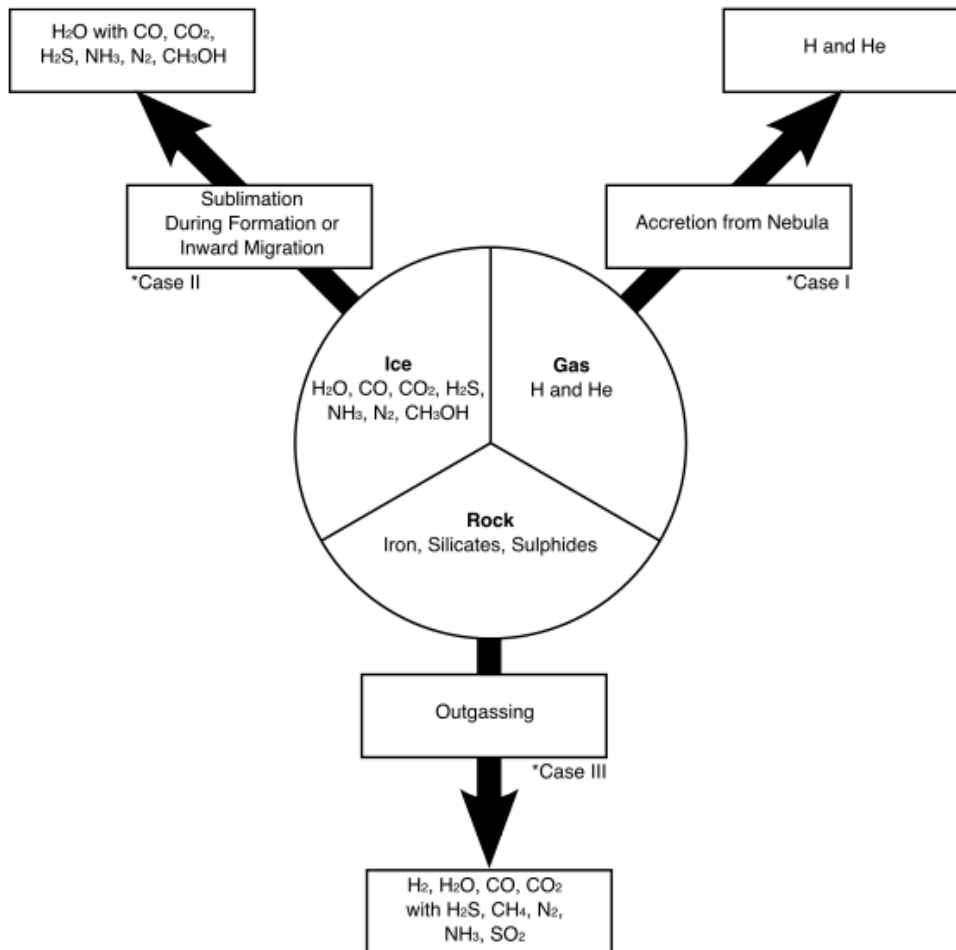
Collisions can significantly alter the planetary compositions (Mercury, Moon + Earth)



The evaporation of the external layers should be significant for close-in planets

Internal structure of super-Earths

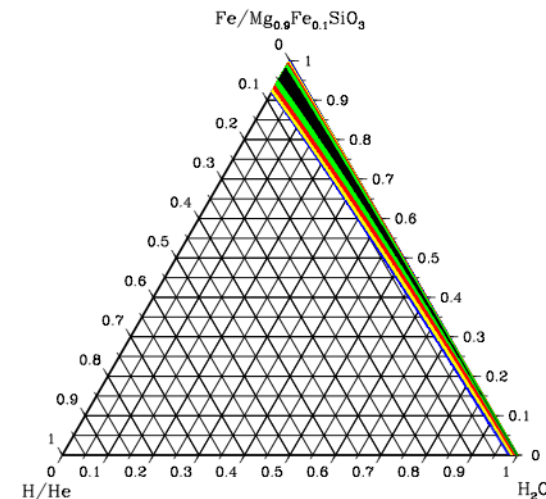
Additional degeneracy: possible massive gas envelope



Rogers & Seager (2010)

Origins ?

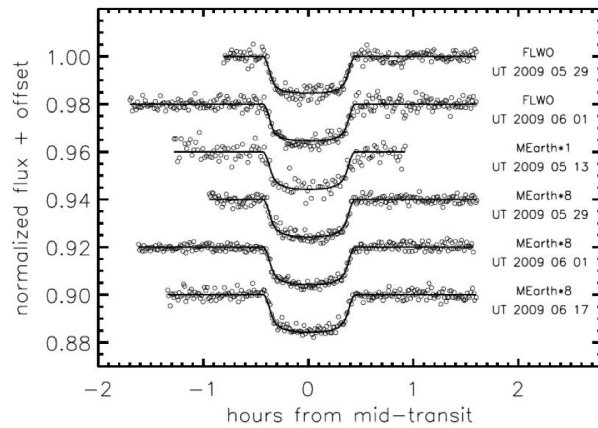
- Outgassing during formation
- Volcanism
- Sublimation of surface layers (ices, rocks for very irradiated planets)
- Accretion of H/He
- Impacts



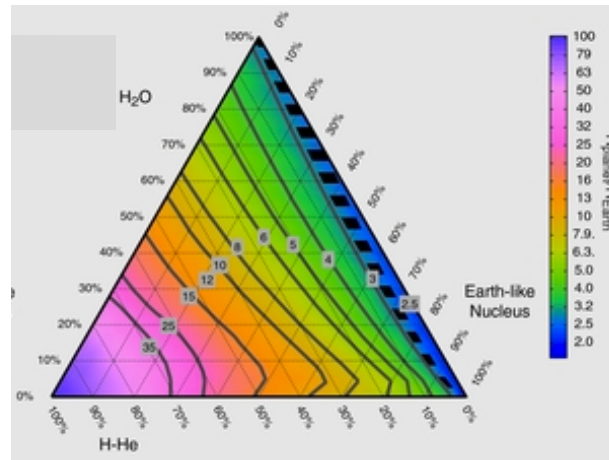
Internal structure of terrestrial and super-Earth planets

Solution: precise measurement of M_p and R_p AND study of the atmosphere

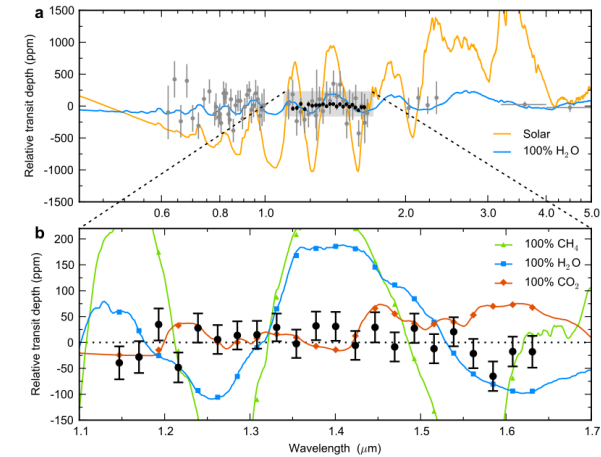
Ex: GJ1214b



Charbonneau et al. (2009)

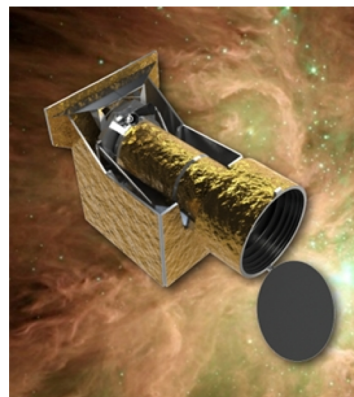


Valencia et al. (2013)



Kreidberg et al. (2014)

Target nearby stars!



CHEOPS



TESS

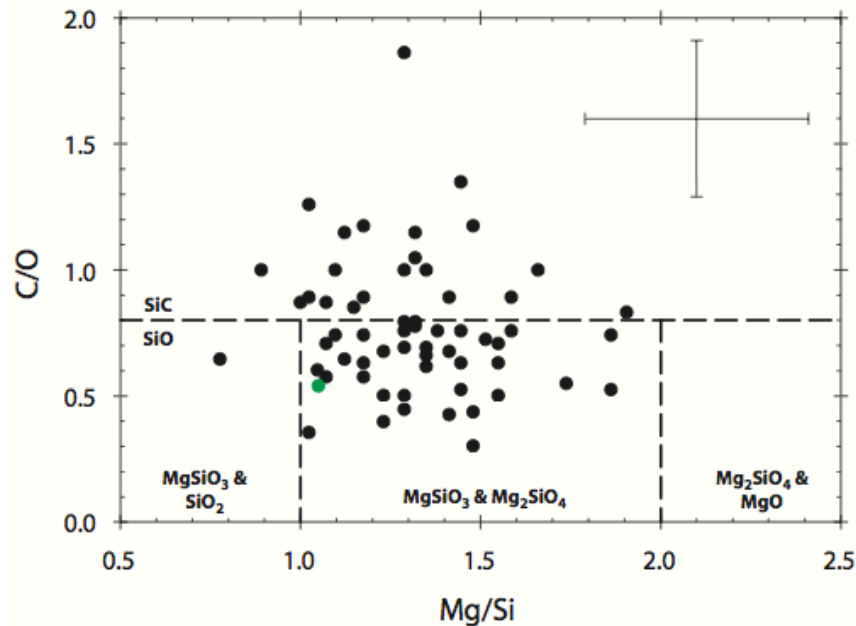


SPECULOOS

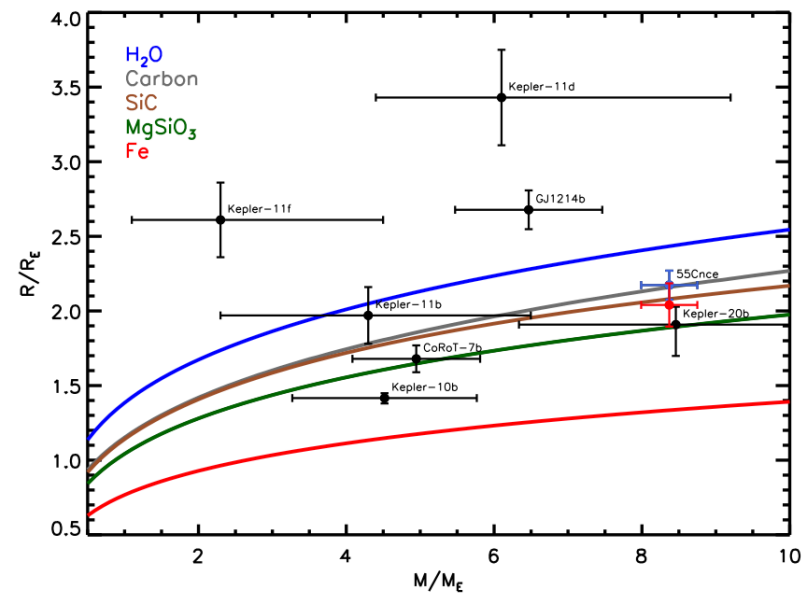
Carbon planets

If C/O doubled compared to solar value in the protoplanetary disk, most silicates would be replaced by silicon carbides and other carbon components

Kuchner & Seager (2007)



Bond et al. (2010)

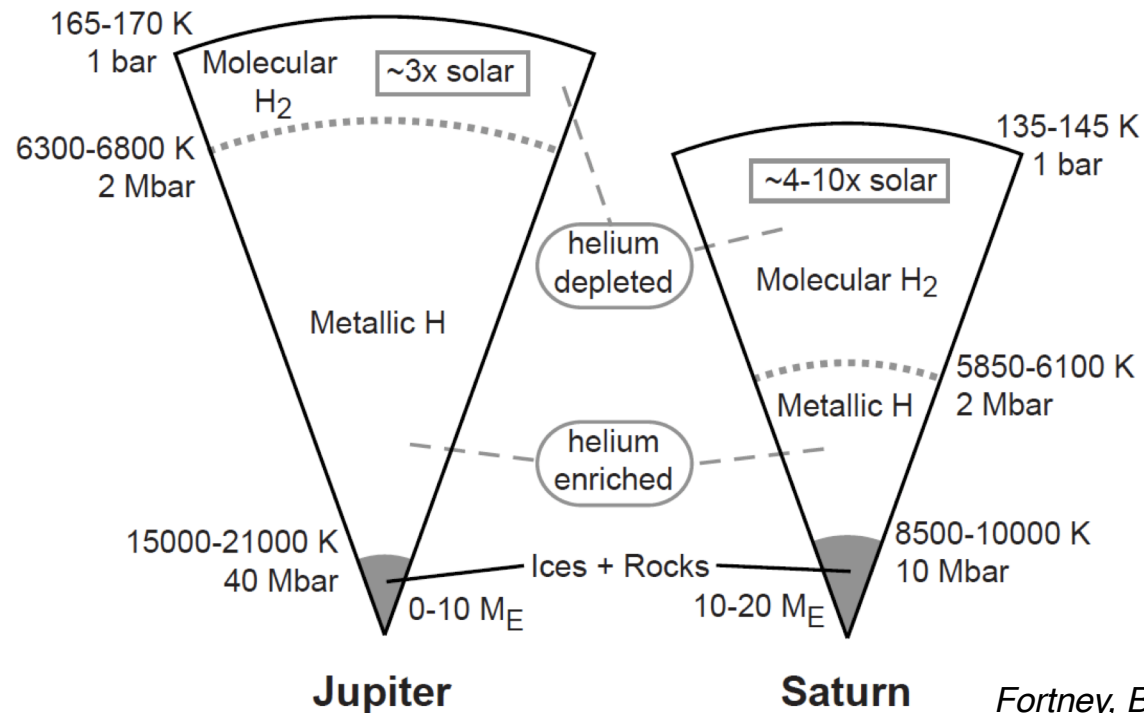


Madhusudhan et al. (2012)

C/O increases with Fe/H and towards the center of the Galaxy

Internal structure of giant planets

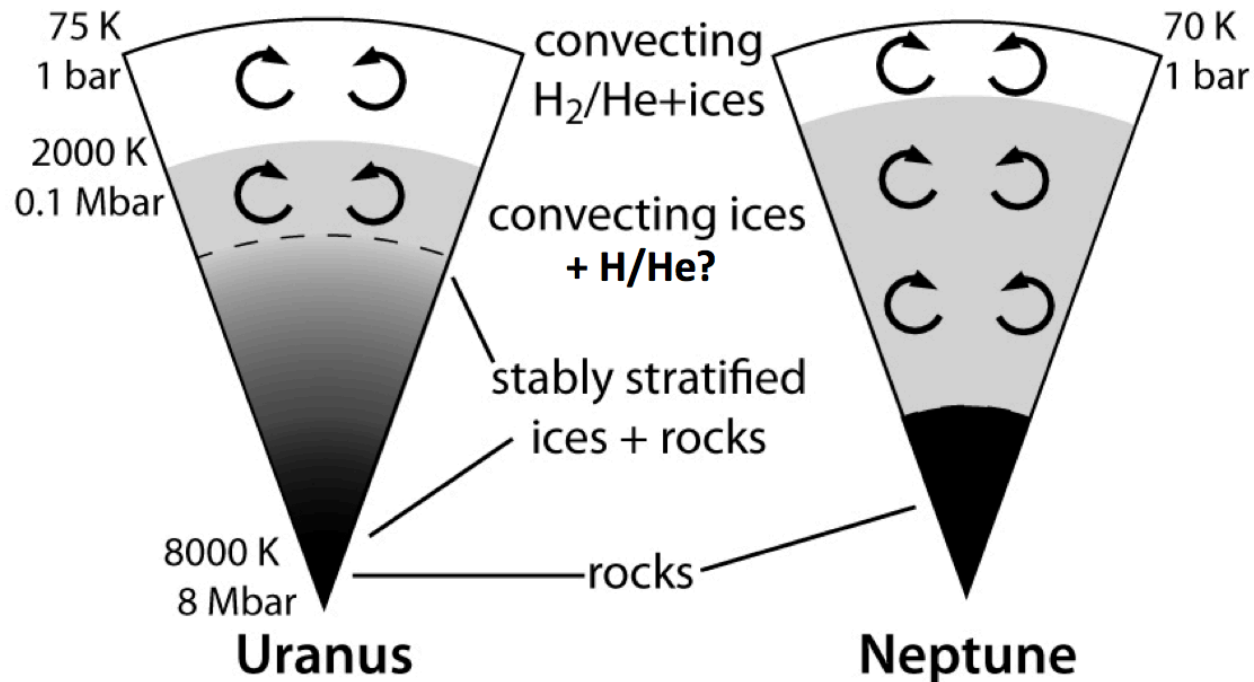
Jupiter and Saturn



Metallic hydrogen: trellis of protons with a spacing \ll Bohr radius (degenerate matter), uncoupled of the electrons \rightarrow highly conductor (heat and electricity). Required pressure: ~ 2 MBar
Coupled to Jupiter's rotation \rightarrow strong magnetic field

Internal structure of giant planets

Uranus and Neptune



Fortney, Baraffe, & Militzer (2010)
"Exoplanets" book, Arizona Space Science Series

External H-He envelope = 5-10% of the mass of the planet, 10–20% of the radius

Internal structure of giant planets

Giant planet model

First approximation : are neglected the magnetic field, the rotation, and the radiative pressure → fluid object supported by gravity and gas pressure

Equations

Same as for a terrestrial planet + **energy conservation equation**

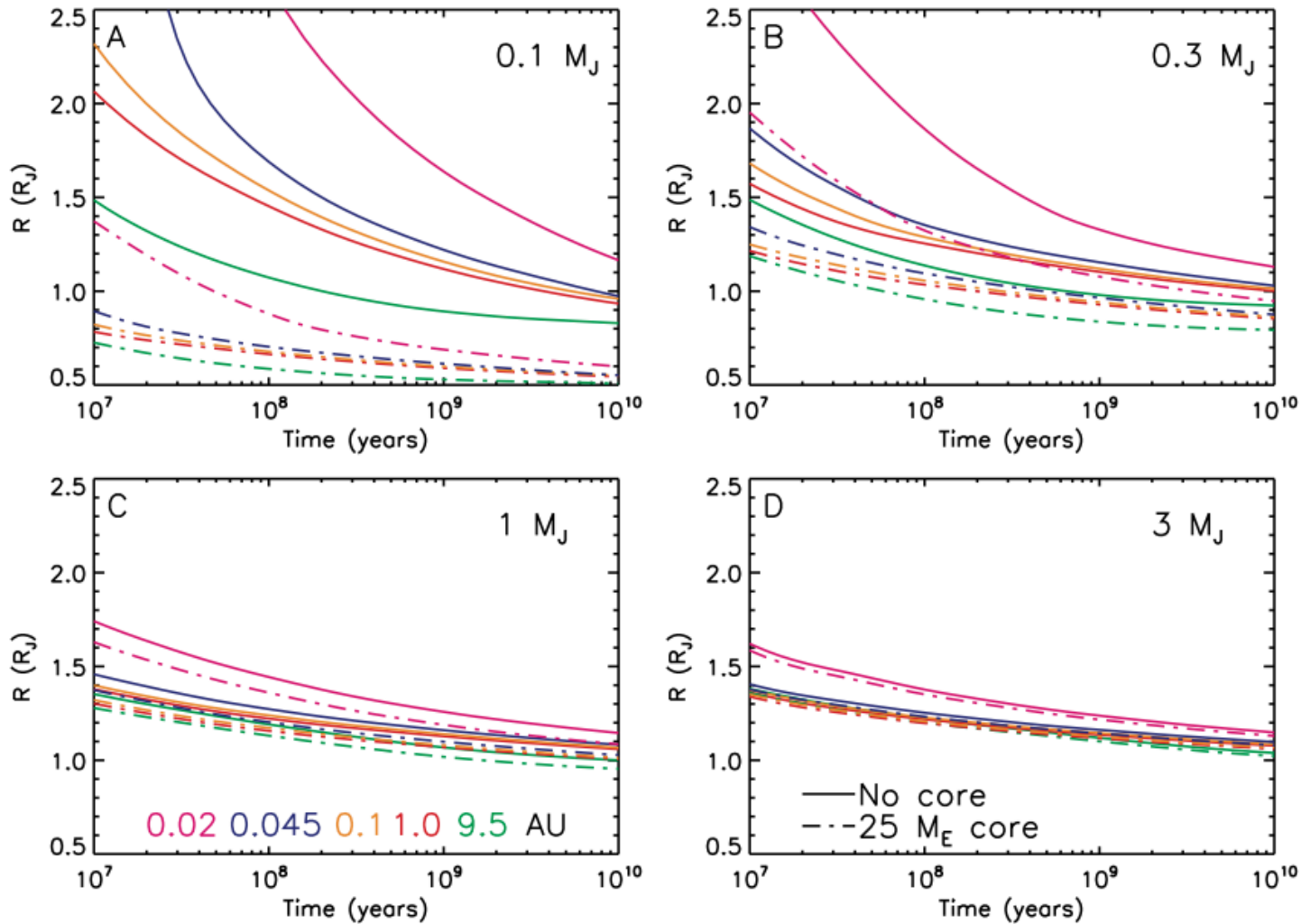
$$\frac{\partial L}{\partial r} = 4\pi r^2 \rho \left(\dot{\epsilon} - T \frac{\partial S}{\partial t} \right)$$

Energy source (planetesimal accretion
+ host star irradiation + tidal heating)

Energy loss (emission)

Internal structure of giant planets

Evolution of the radius of an irradiated giant planet



Fortney et al. (2007)

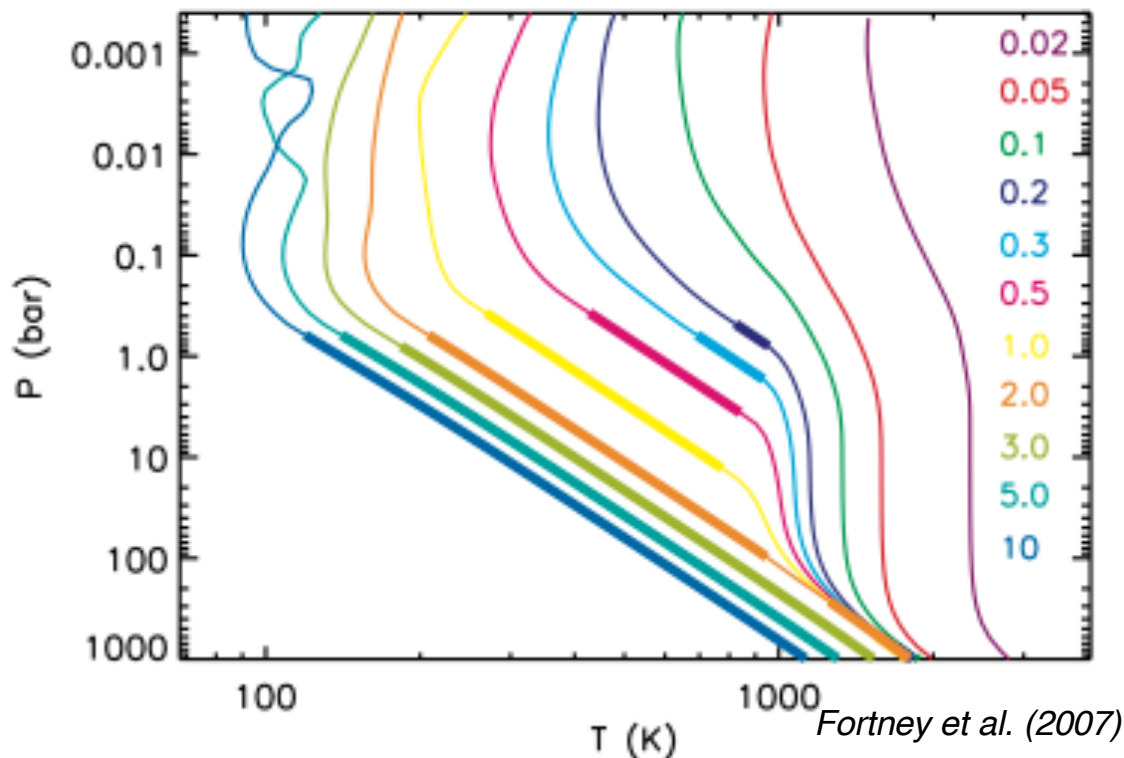
Internal structure of giant planets

Thermal evolution and connection interior-atmosphere

The atmosphere is the bottleneck that controls the deposit (irradiation) and the release of the energy.



The structure model must be coupled to an atmospheric model

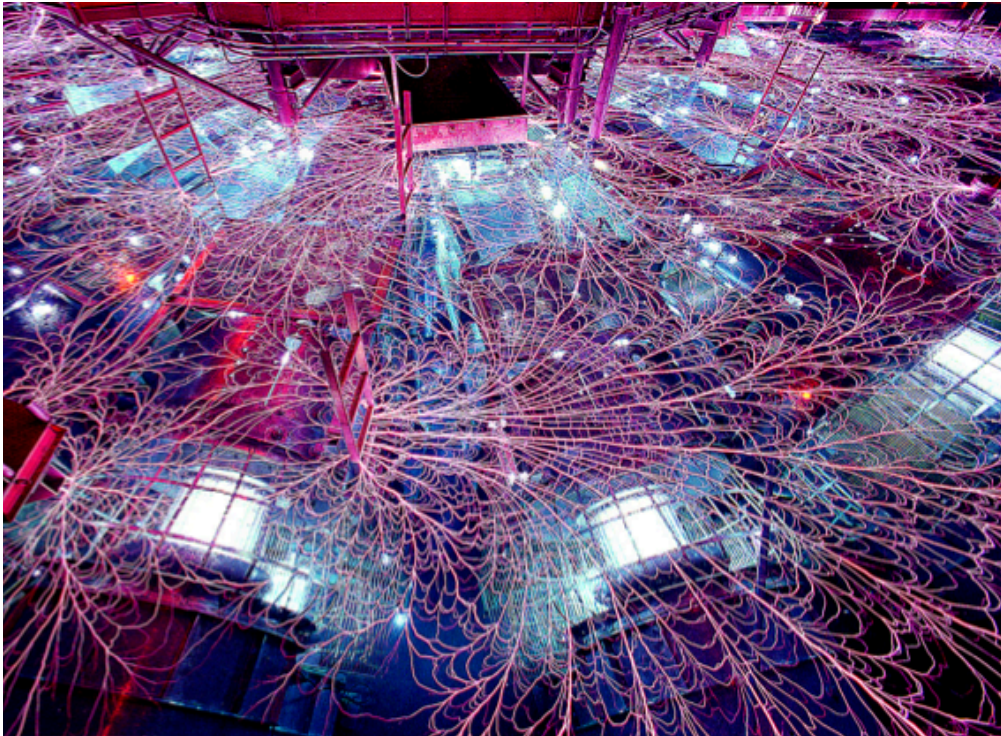


P-T profiles for
4.5 Gy Jupiter-
like planets

Thick line =
convective
regions

Internal structure of giant planets

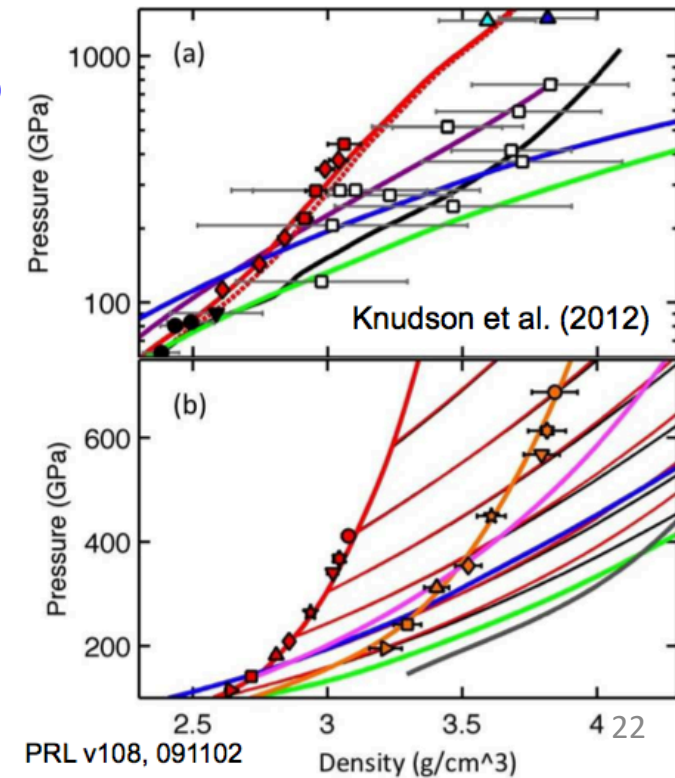
Equation of state at very high pressure and temperature



Sandia National Laboratory Z Machine
High P ($>1\text{TPa}$) and T ($>2\text{GK}$) reached
through high electromagnetic fields

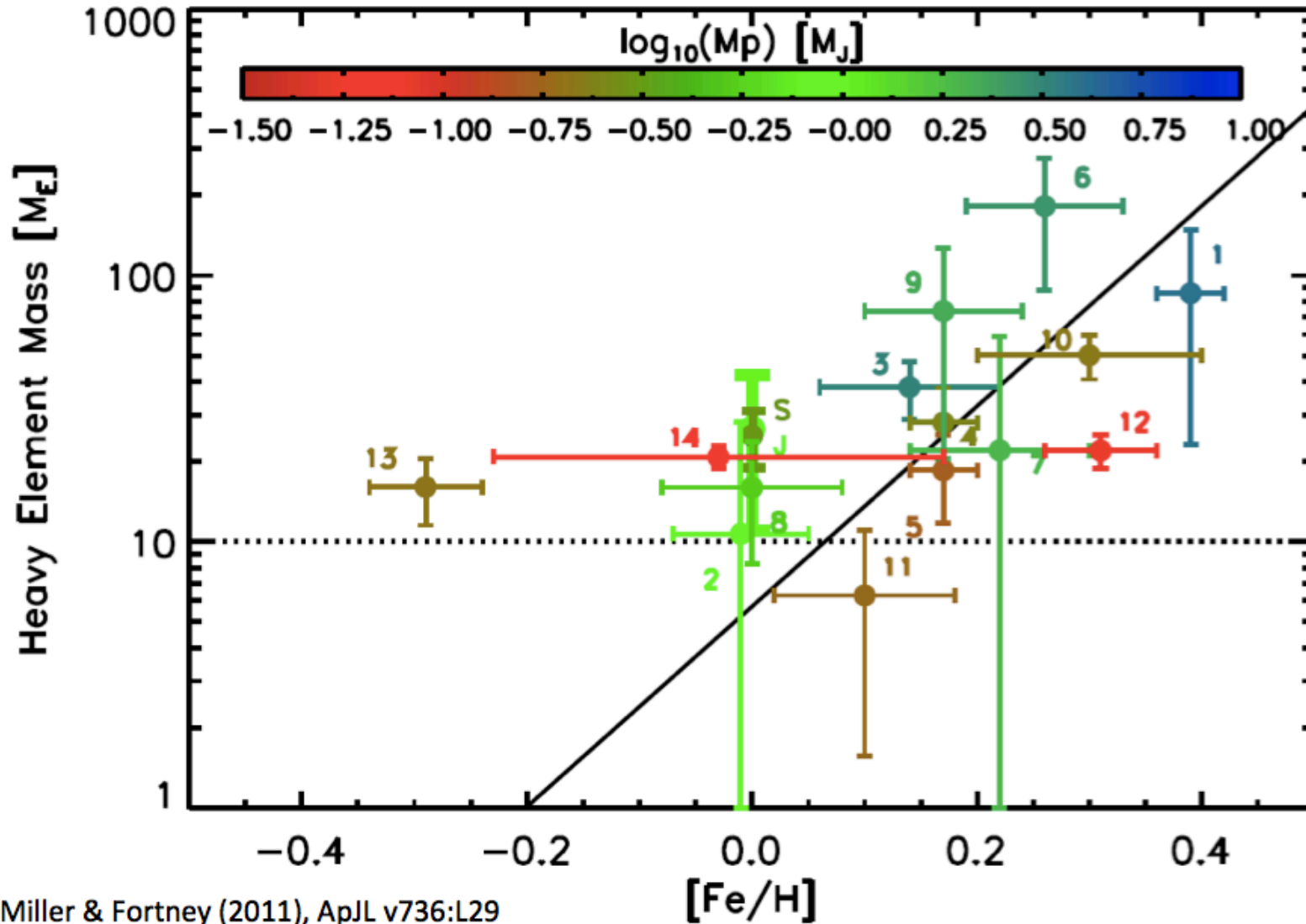
Recent significant advances
via
Ab initio computations
Precise measurements

H_2O



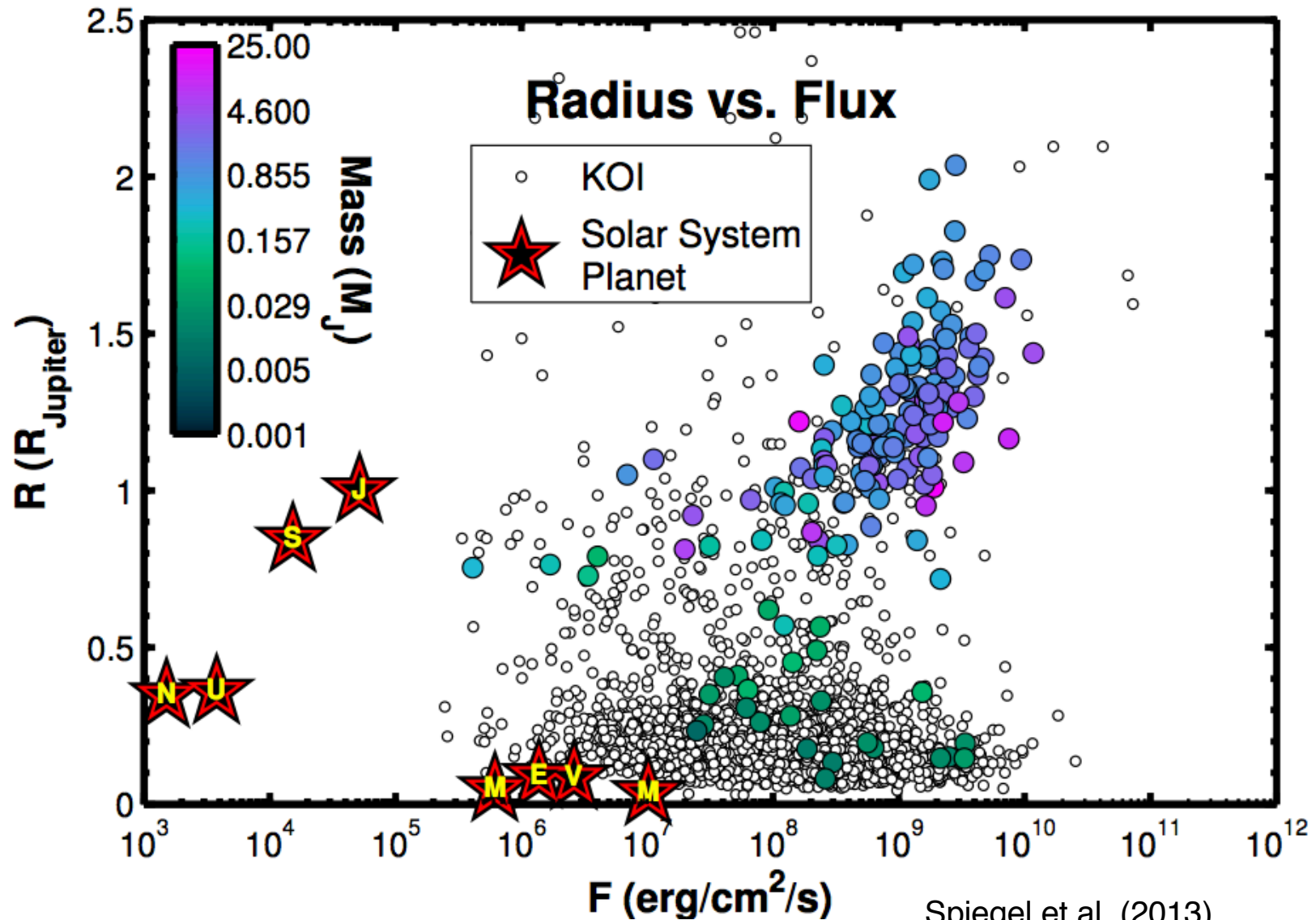
Internal structure of giant planets: results

Correlation between stellar and planetary abundances?



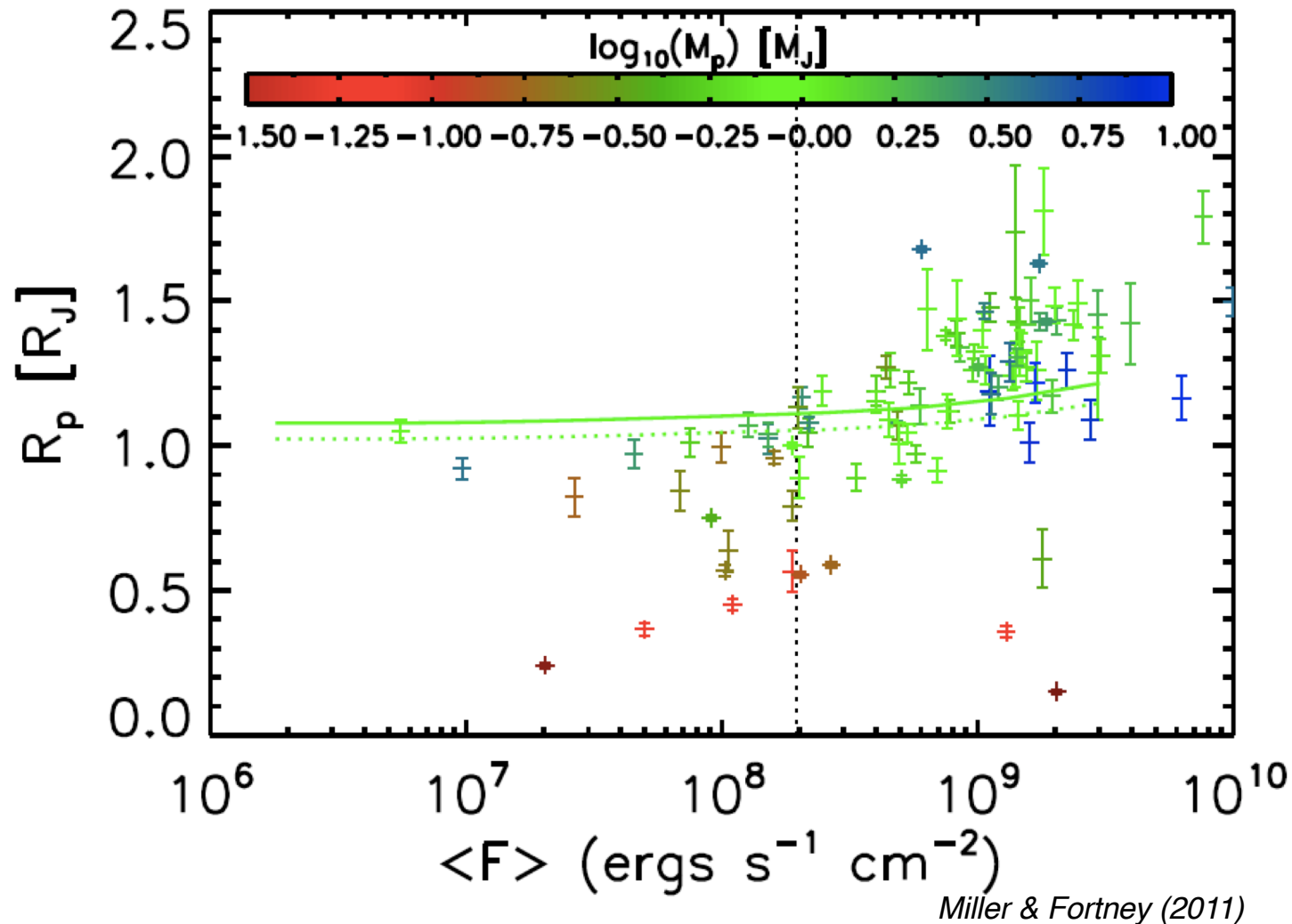
Internal structure of giant planets: results

Relationship radius – irradiation for giant planets



Internal structure of giant planets: results

The most irradiated giant planets are often 'too' big



Internal structure of giant planets

Origins of that “radius anomaly” ?

1/ Deep deposit (at the radiative-convective limit, or lower) of 0.1 to 1% of incident irradiation

Mechanism? Day-night winds (*Guillot & Showman 2002*); ohmic dissipation (*Batygin & Stevenson 2010*)

2/ Slowed down contraction

Current theories predict contractions too fast to explain radius anomalies.

Additional mechanism? Larger opacities (*Burrows et al. 2007*); convection in layer due to composition gradient (*Chabrier & Baraffe 2007*)

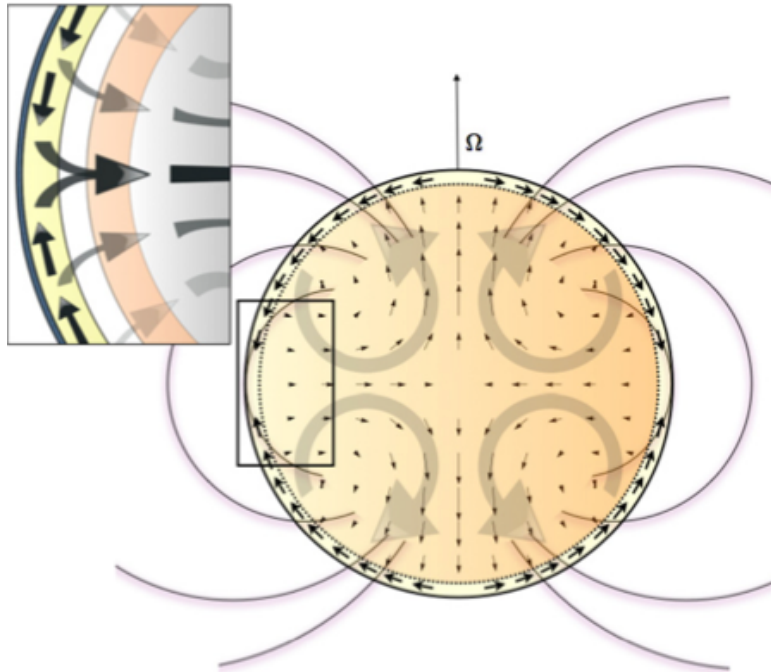
3/ Tidal dissipation

Stellar tides: strong effect but drops to zero once the orbit is fully circularized (*Bodenheimer et al. 2001*)

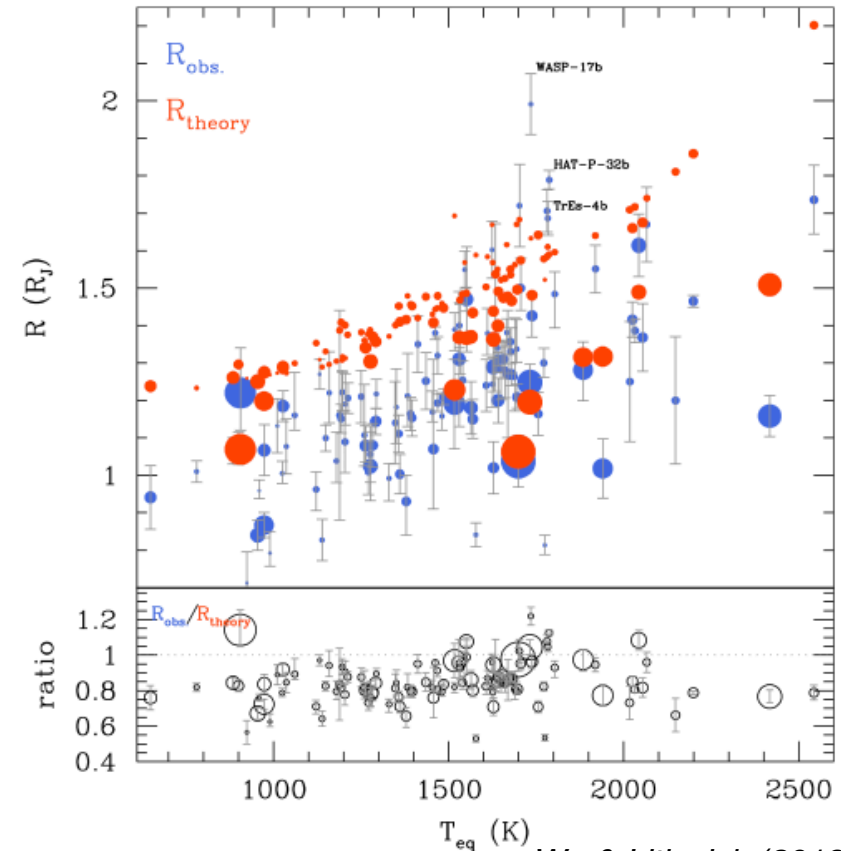
Thermal tides? *Arras & Socrates (2010)*

Internal structure of giant planets

Ohmic dissipation



Batygin & Stevenson (2010)



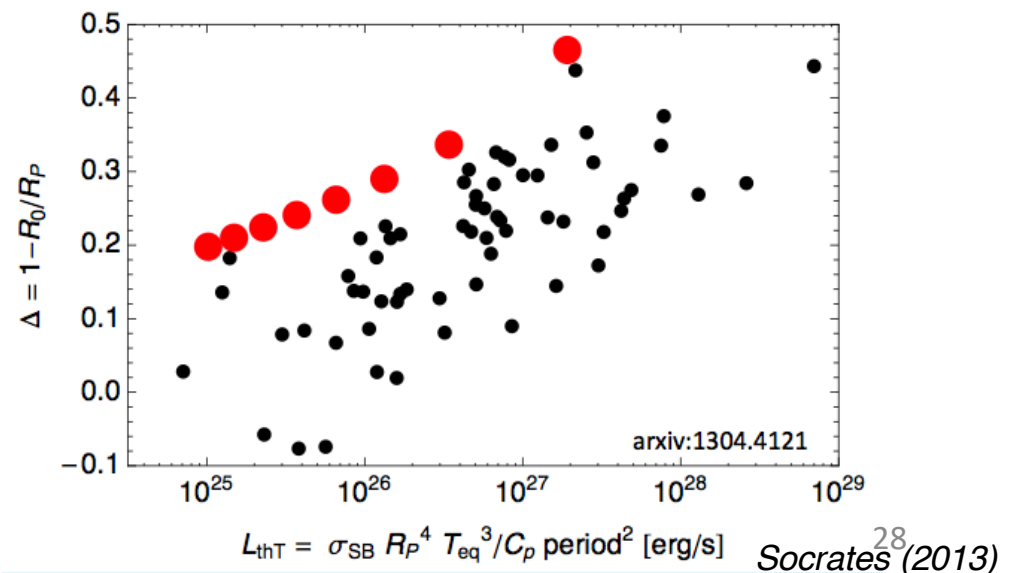
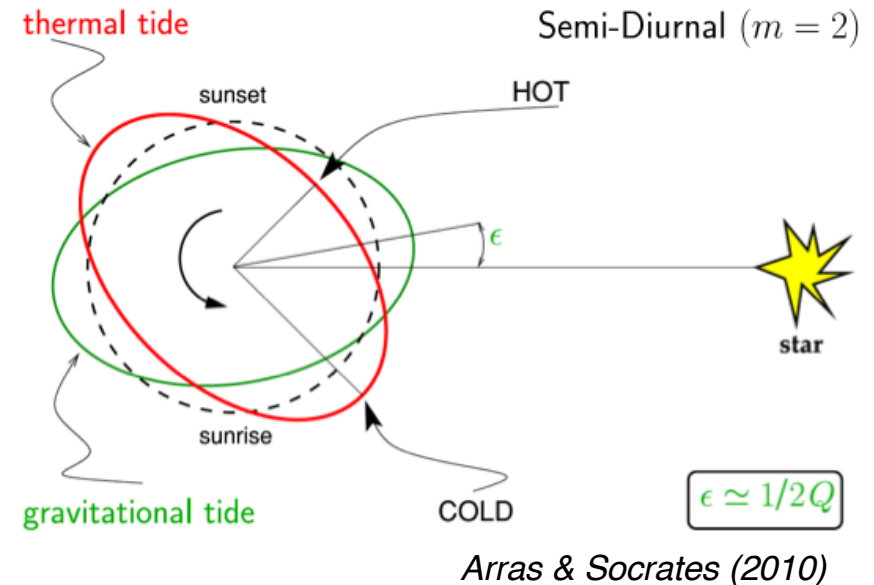
Wu & Lithwick (2013)

- Alkaline metals (Li, Na, K,...) thermally ionized and carried by advection in the strong winds of the external layer.
- Creation of a magnetic field distinct from the inner magnetic field (dynamo)
- Induced current deposits its energy by ohmic dissipation in the internal layers

Internal structure of giant planets

Thermal tides

- ◆ The thermal inertia of the atmosphere is generally not negligible
- ◆ The peak of T (afternoon) happens later than the irradiation peak (noon)
- ◆ The attraction of the star on the thermal bulge tends to make the planet spin faster
- ◆ Opposition between gravitational and thermal tides \rightarrow constant energy source
- ◆ Outcome: a tidal mechanism with irradiation as energy source

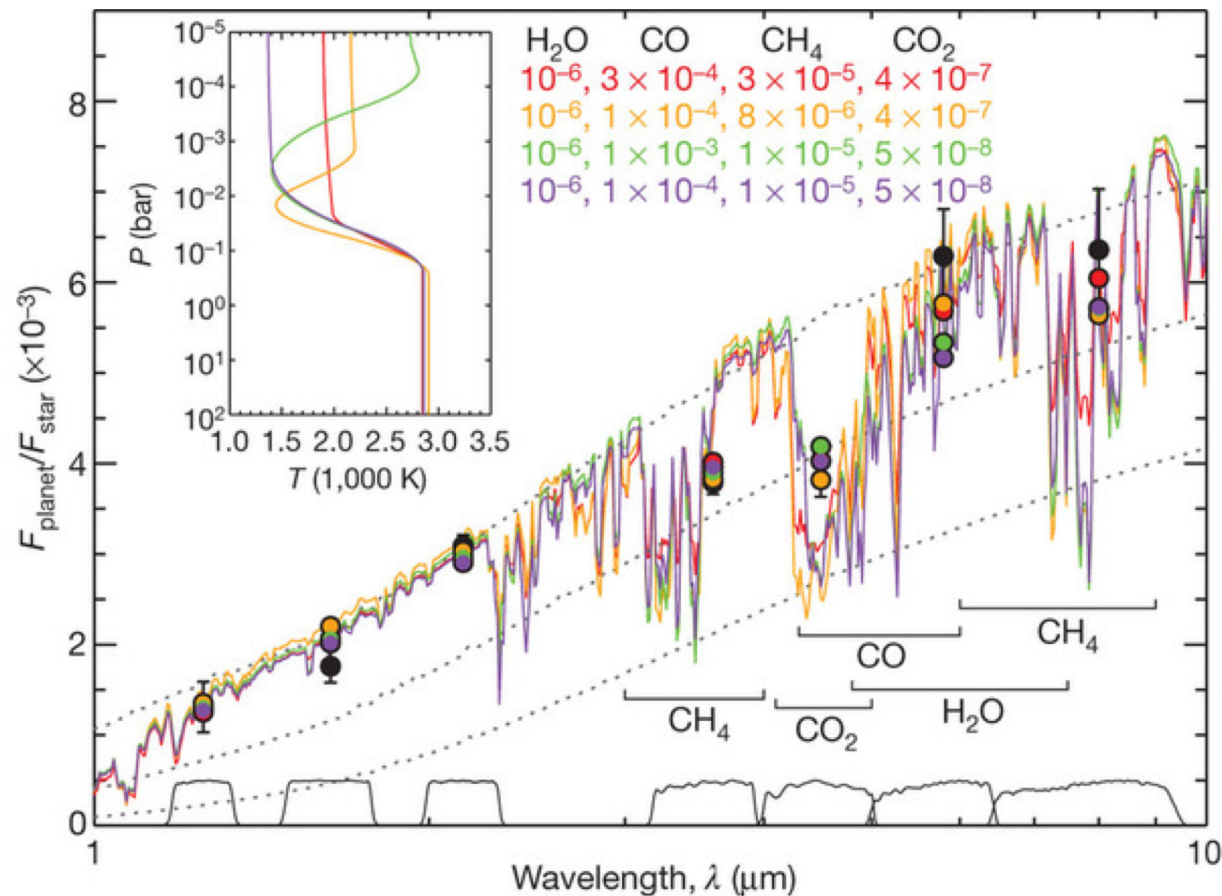


Planetary atmosphere

Study of exoplanet atmospheres

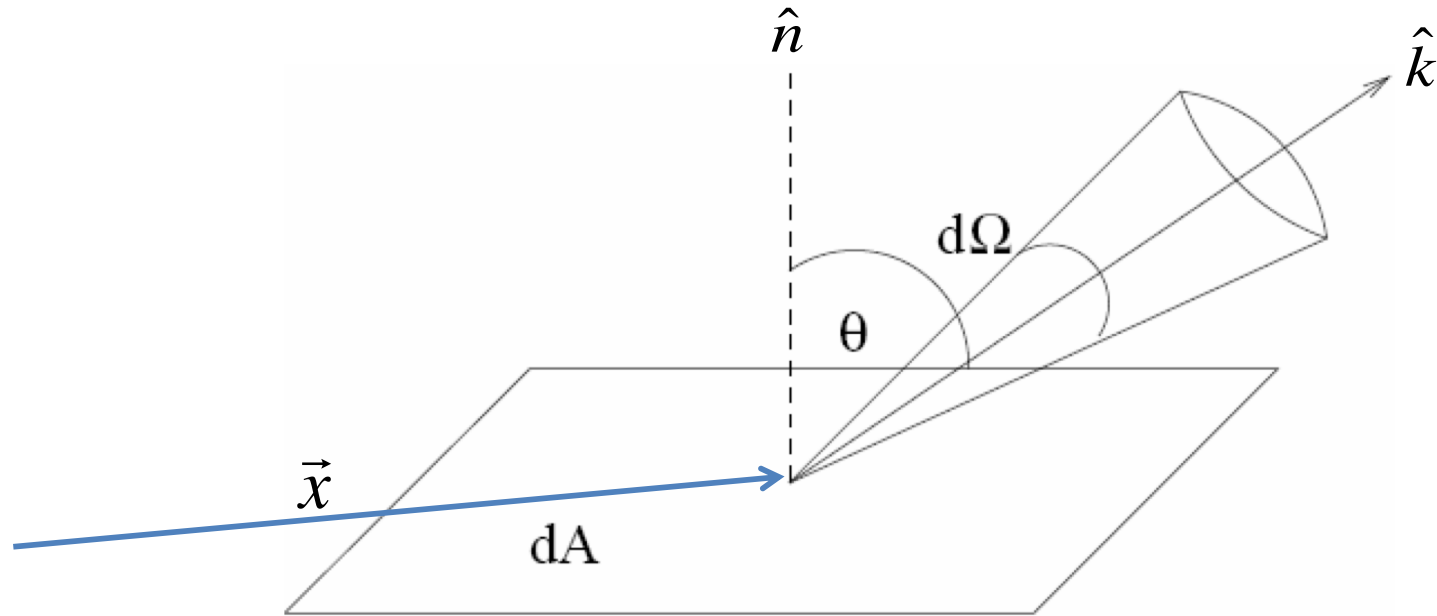
Composition, circulation, clouds, thermal structure, etc.

→ spectroscopy then inferences based on the comparison with radiative transfer models



Planetary atmosphere

Spectral radiance (a.k.a. specific intensity)



$$dE_\nu = I_\nu \cos \theta dA d\Omega dt d\nu$$

Spectral radiance $I_\nu = I(\nu, t, \vec{x}, \hat{k})$

Units: $\text{W m}^{-2} \text{sr}^{-1} \text{Hz}^{-1}$

Planetary atmosphere

Flux

$$\vec{F}_\nu(\vec{x}, \nu, t) = \int_{4\pi} I_\nu(\vec{x}, \nu, t, \hat{n}) \hat{n} d\Omega$$

Measurement of the quantity of energy crossing dA at a given time t and at a frequency ν

We are generally interested in the flux
in our direction

$$F_\nu = \int_{4\pi} I_\nu \hat{n} \cdot \hat{k} d\Omega$$

$$F_\nu = \int_{4\pi} I_\nu \cos \theta d\Omega$$

Planetary atmosphere

We have thus $dE_\nu = I_\nu \cos \theta dA d\Omega dv dt$



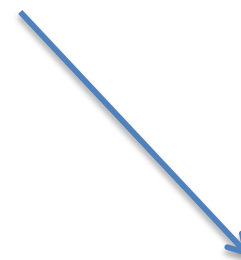
Spectral radiance = the energy in a beam at a given time

The equation of radiative transfer describes how this energy varies due to the interactions between light and matter → variations of I_ν as a function of the length of the path travelled in the medium

$$dI_\nu = -\kappa_\nu \rho I_\nu ds + j_\nu \rho ds$$



Absorption cross section



Emission cross section

ds = infinitesimal distance travelled

ρ = density of the crossed medium

Planetary atmosphere

The equation of radiative transfer

$$\frac{dI_\nu}{\kappa_\nu \rho ds} = -I_\nu + S_\nu$$

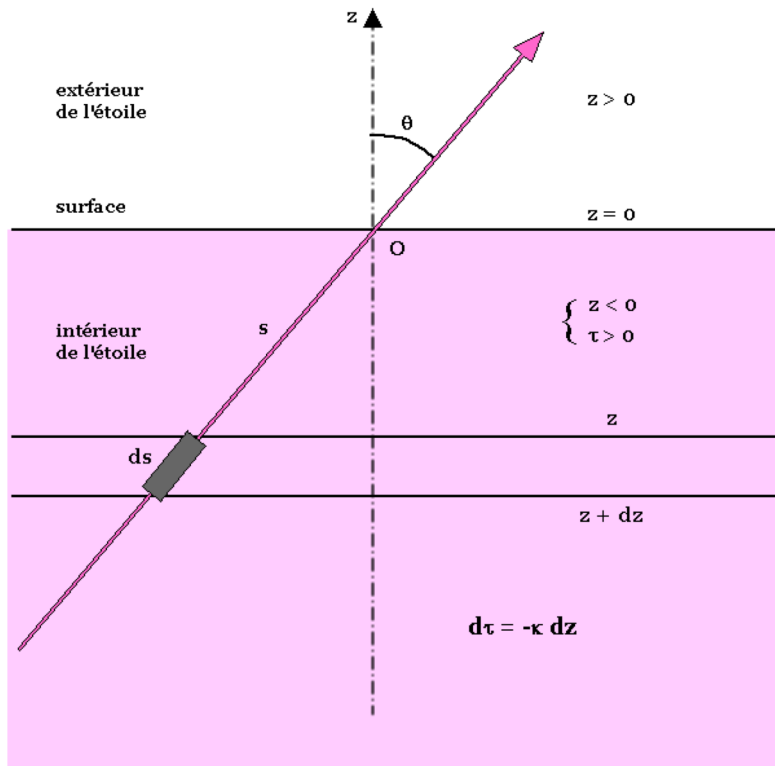
With $S_\nu = \frac{j_\nu}{\kappa_\nu}$ the source function

Apparently simple equation, but

- Not linear: I_ν depends on j_ν , but j_ν depends also on I_ν
- Opacities behind κ_ν and j_ν depend on millions of atomic and molecular lines, and in the case of opacity due to clouds, include many free parameters.

Planetary atmosphere

Plane-parallel atmosphere



The curvature of the atmospheric layers is neglected

If
$$\mu = \cos \theta$$

Then
$$dz = \cos \theta ds$$

Is then defined the optical depth τ_v , measured from the top of the atmosphere, as:

$$\tau_v(z) = - \int_z \kappa_v(z) \rho(z) dz$$

→ Equation of radiative transfer becomes

$$\mu \frac{dI_v}{d\tau} = I_v - S_v$$

Planetary atmosphere

Analytical solutions. I. Transmission in the optical

Thermal emission of the atmosphere negligible: $S_\nu = 0$

$$\mu \frac{dI_\nu}{d\tau} = -I_\nu \quad \longrightarrow \quad I_\nu = I_\nu(0) e^{-\tau_\nu / \mu}$$

Beer-Lambert law

II. Thermal emission without scattering, local thermodynamical equilibrium (LTE)

$$S_\nu = B_\nu \quad \text{Source function} = \text{Planck function}$$

$$\mu \frac{dI_\nu}{d\tau} = I_\nu - B_\nu \quad \longrightarrow \quad I_\nu(z) = \int_0^\pi \frac{1}{\mu} \int_0^\infty B_\nu(\tau) e^{-\tau_\nu(z)/\mu} d\tau d\mu$$

Planetary atmosphere

Chemical composition

An elementary composition is adopted (e.g. solar)

Assumed network of chemical reactions

The chemical equilibrium is assumed (no change of abundances)

The abundance of a given molecule in a given layer depends on P and T (in first approach).

Computation of abundances through **minimization of Gibbs free energy G**

First law of thermodynamic: $dU = dQ - dW = TdS - PdV$

Definition of G: $G = U + PV - TS$

$$dG = dU + PdV + VdP - SdT - TdS$$

$$dG = VdP - SdT$$

$dG \leq 0$ until equilibrium for which $dG=0$

Planetary atmosphere

Chemical composition

Determination of the minimum of

$$G/RT = \sum_i n_i \left[\left(g_i^0(T)/RT \right) + \ln P + \ln(n_i/N) \right]$$

with

$$\sum_i a_{ij} n_i = b_j$$

n_i = number of moles of molecule i

g_i^0 = Gibbs energy of molecule i at standard pressure (**Tables**)

N = total number of moles of the system

b_j = number of moles of element j

a_{ij} = number of atoms of element j in molecule i

A charge constraint can also be used to take into account ions

Solid particules? Same but no pressure term

Planetary atmosphere

Departure from chemical equilibrium

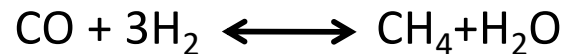
If P and T are high, chemical reactions are fast and equilibrium is quickly reached.

But in the cold low-density layers, it can never be reached...

e.g. T-type brown dwarfs ($T_{\text{eff}} < 1300\text{K}$)

CO is the dominant form of carbon-bearing in the deep layers, and CH₄ should totally dominate the upper colder layers. Still, they also contain CO.

It is due to this following reaction which is there very slow:

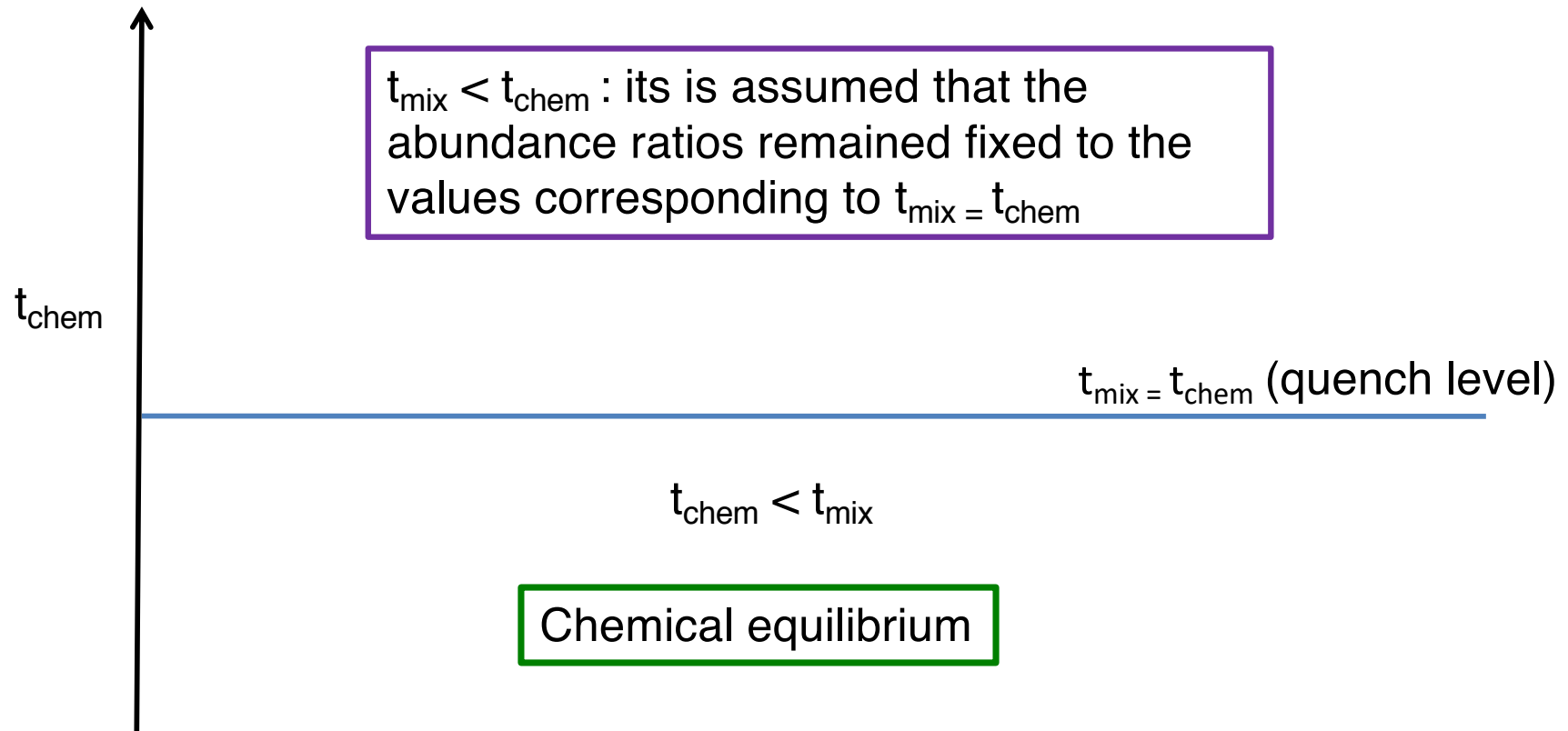


The vertical mixing rate of CO is much superior to its rate of transformation into CH₄, so CO accumulates in the upper layers.

Same thing happens in Jupiter's atmosphere.

Planetary atmosphere

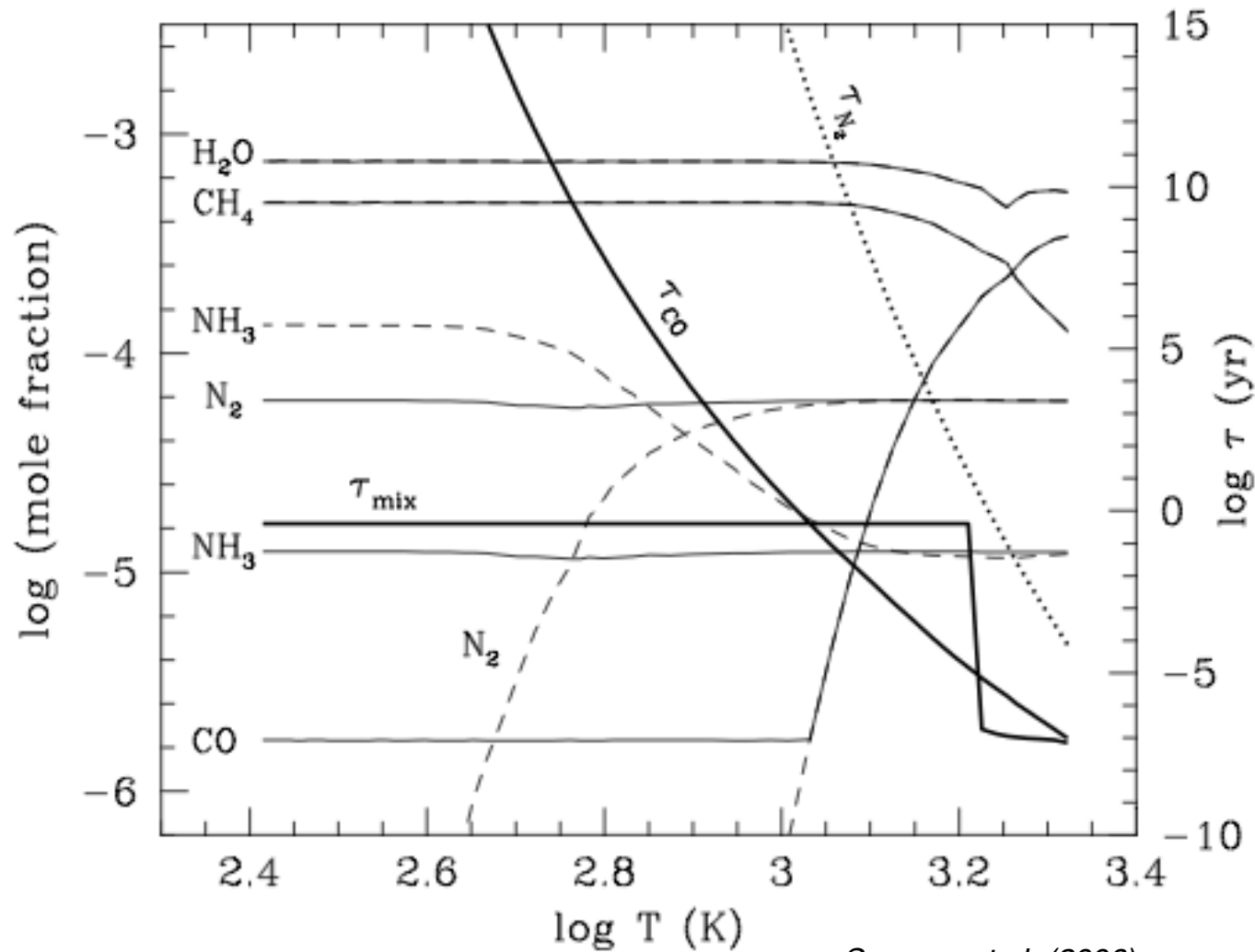
Departure from chemical equilibrium: modeling



t_{mix} is generally poorly constrained...

Planetary atmosphere

Departure from chemical equilibrium



Saumon et al. (2006)

Planetary atmosphere

Atmospheric model (simplest version)

Code of radiative transfer (1D plane-parallel + LTE).

Resolution for each layer of the equation of radiative transfer

+ thermodynamic equilibrium (radiation or convection-dominated heat transfer) + hydrostatic equilibrium, so to obtain $P(z)$, $T(z)$, and $I(z,v) =$

atmospheric profile

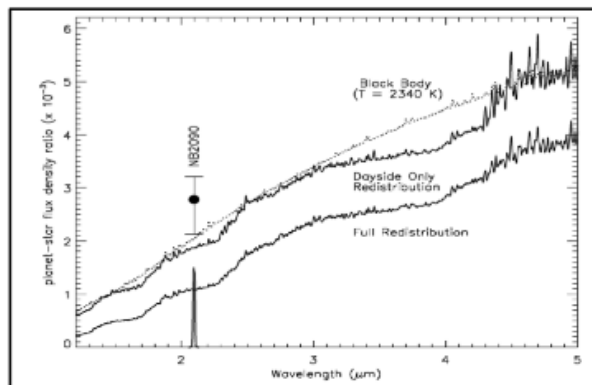
Limit conditions: stellar irradiation and internal entropy

Parameters:

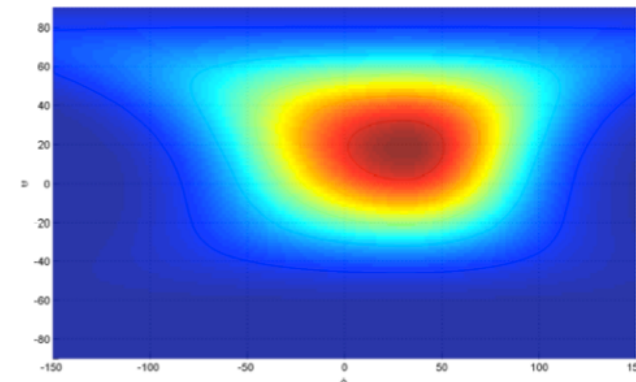
- Internal entropy
- Irradiation
- Opacities
- Choice of atomic and molecular species
- Elementary composition
- Chemical equilibrium or not (vertical mixing)
- Clouds (opacities)
- Efficiency of the day-night heat transfer

Planetary atmosphere

1D vs CGM 3D models



Gillon et al. (2010)



de Witt et al. (2012)

Habitability

Habitability criteria (NASA)

1. *Extended areas of liquid water*
2. *Conditions suitable to the assembly of complex organic molecules*
3. *Sources of energy available to sustain metabolisms*



Exoplanets

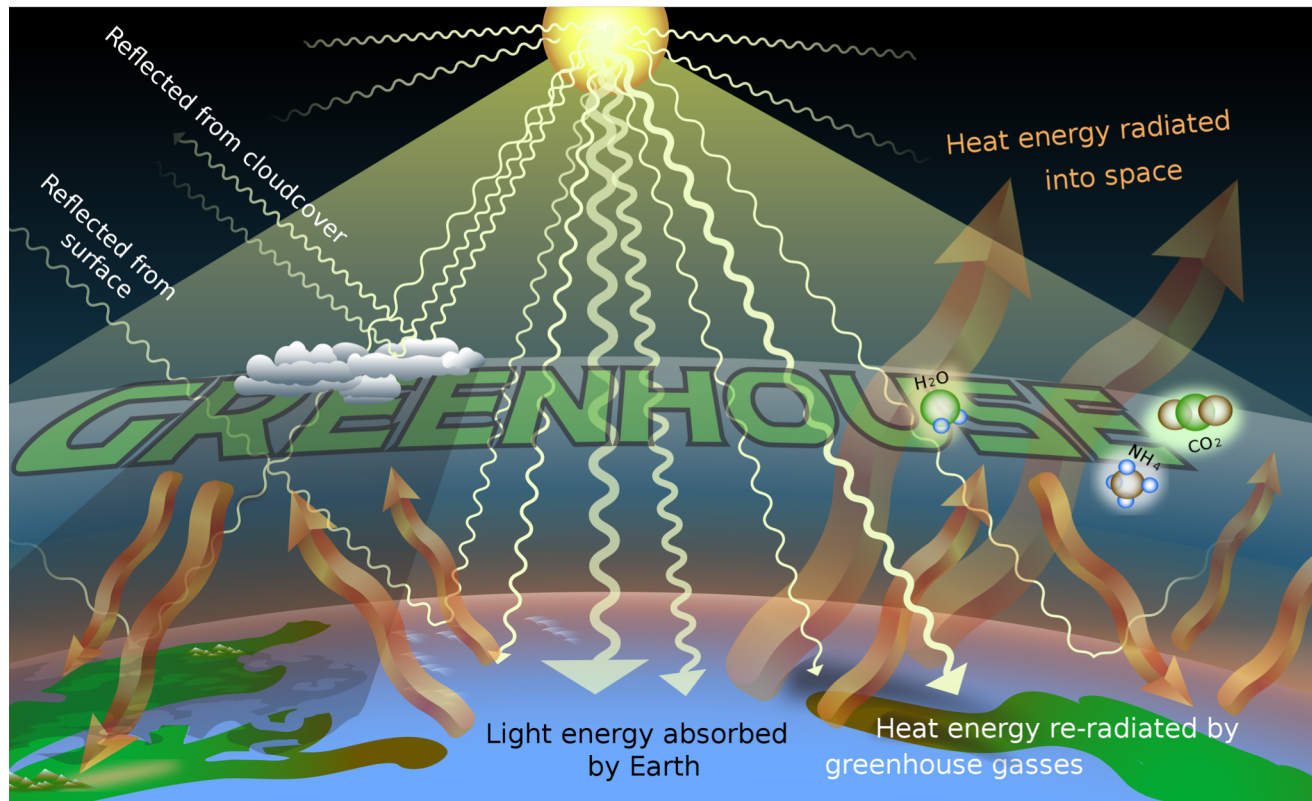
Focus is on terrestrial planets that could harbor liquid water on their surface



Notion of Habitable Zone (HZ)

The greenhouse effect

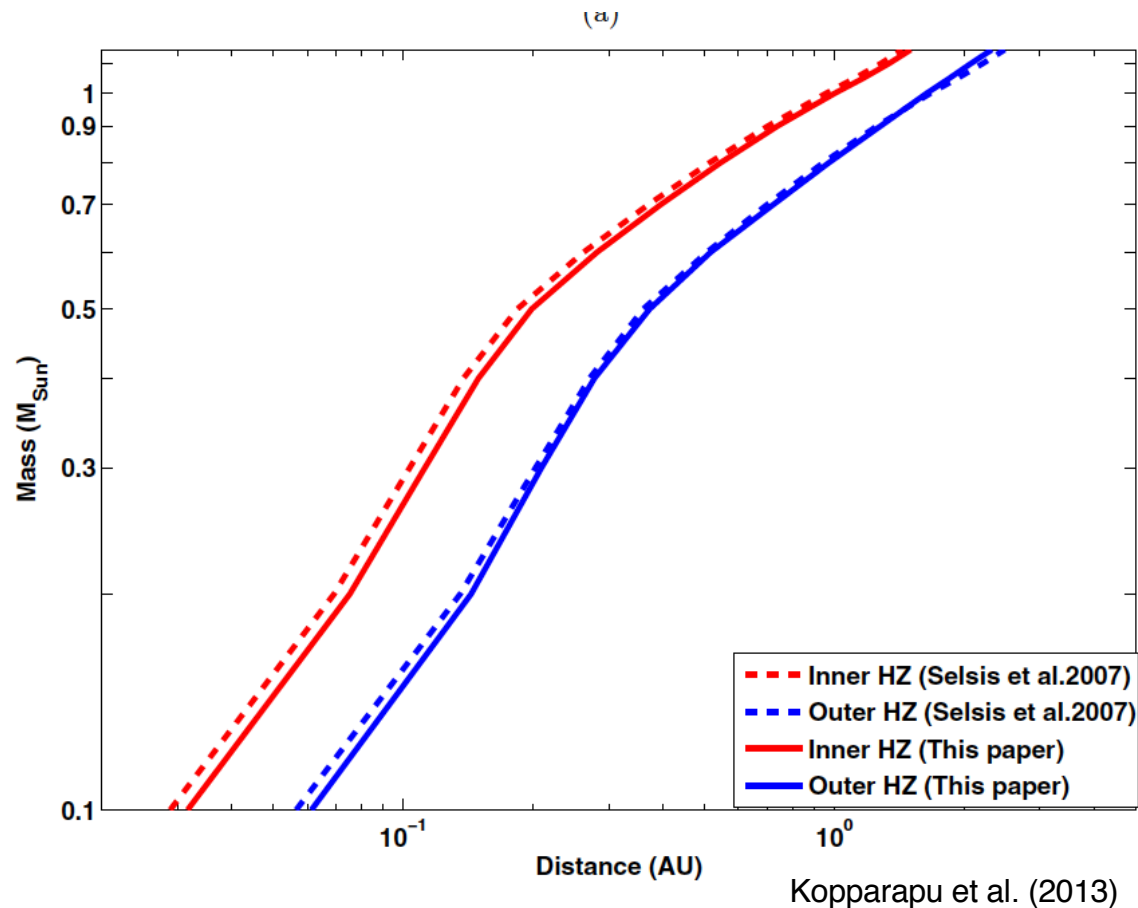
Trapping of radiation within the atmosphere by absorption & isotropic reemission of the infrared emission from the surface or lower atmospheric layers



Earth's greenhouse gases: $\text{H}_2\text{O} > \text{CO}_2 > \text{CH}_4 > \text{N}_2\text{O} > \text{O}_3 > \text{CFCs}$

The habitable zone

Distance range to a star in which a terrestrial planet with suitable atmospheric features could sustain liquid water on its surface (Huang 1959)



Inner limit:

Runaway greenhouse state \rightarrow water loss by photolysis and hydrogen loss

Outer limit:

Formation of CO_2 clouds that cool down the surface by increasing the albedo and decreasing the greenhouse effect

Solar system: from 0.99 AU to 1.7 AU (Kopparapu et al. 2013)

The greenhouse effect: importance of CO₂

Main CO₂ sources:

1) **Decarbonation of carbon-containing rocks in subduction zones:**

silicates + CaCO₃ → other silicates + CO₂ → Outgassed by volcanoes, hot springs, etc.

2) **Uplifting of carbon-rich sediments** (mostly of biological source) during mountain building (or by human intervention), **which are oxidized and release CO₂.**

3) **Decomposition/respiration:** (CH₂O)_n + nO₂ → CO₂ + H₂O

CO₂ sinks:

1) **Chemical weathering** CO₂ + H₂O (rain water) : H₂CO₃

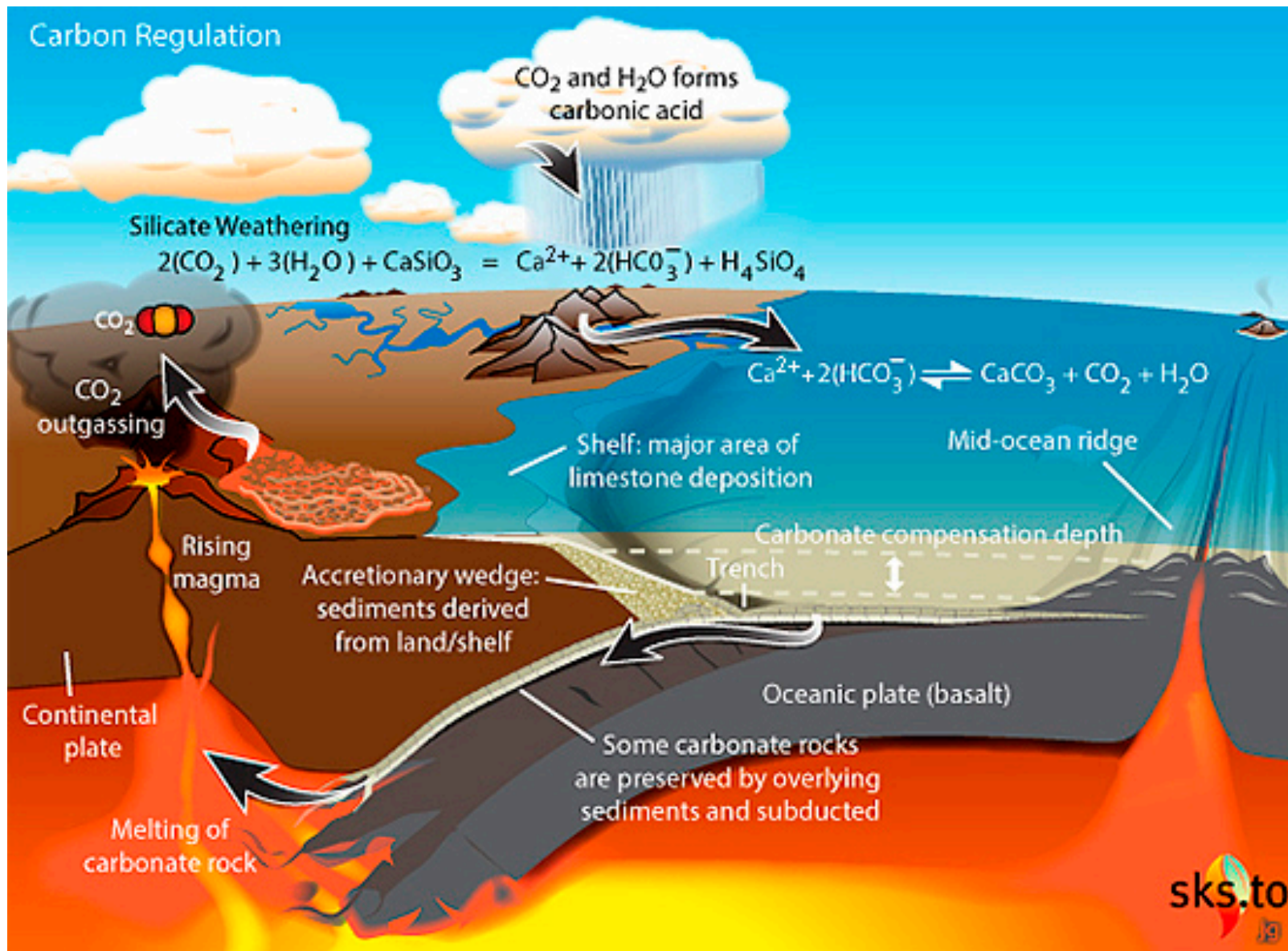
2H₂CO₃ + H₂O + CaSiO₃ → Ca²⁺ + 2HCO₃⁻ + H₄SiO₄ (chemical weathering of silicates)



2) **Burial of carbon-rich sediments** (e.g. late Carboniferous Period)

3) **Photosynthesis:** nCO₂ + nH₂O → (CH₂O)_n + nO₂

The greenhouse effect: importance of CO₂



CO₂ and H₂O: climate feedbacks

CO₂:

If T increases -> more evaporated water -> more chemical weathering
-> less CO₂ in the atmosphere
-> smaller greenhouse effect -> **T decreases**

Negative feedback loop (stabilizing)

H₂O:

If T increases -> more evaporated water -> larger greenhouse effect -> **T increases**
-> More clouds -> **T increases or decreases**

-> less ice -> lower albedo -> **T increases**

Positive feedback loop (destabilizing)

Habitability: a complex multi-factors problem

Parameters

- Axial tilt (importance of the Moon)
- Orbital eccentricity
- Continents' position
- Tides (heating, orbital evolution, tidal locking)
- Atmospheric composition (e.g. H₂ is a strong greenhouse gas)
- Life itself!
- Surface gravity
- Plate tectonism (could be more or less likely for super-Earths)
- Radioactive decay
- Atmospheric erosion
- Volcanism
- Impact rate
- Magnetosphere
- Heat redistribution: oceanic and atmospheric currents, rotation period
- Low insolation regime: regular snowball states
- Spectral type
- Etc...

The galactic habitable zone

The galactic habitable zone

Galactic habitable zone
In this region, rocky planets can form, but nearby supernovas are rare.

“Similar to the Goldilocks zone but on a galactic scale”

Danger zone
Too close to the galactic hub, life would be harmed by deadly rays from frequent stellar explosions.

Outer regions
Far from the galactic hub, galaxies lack the heavy elements needed for life-sustaining rocky planets.

Mars orbit

Mercury orbit

Too hot
Too close to a star, and a planet's surface water and atmosphere boil away into space.

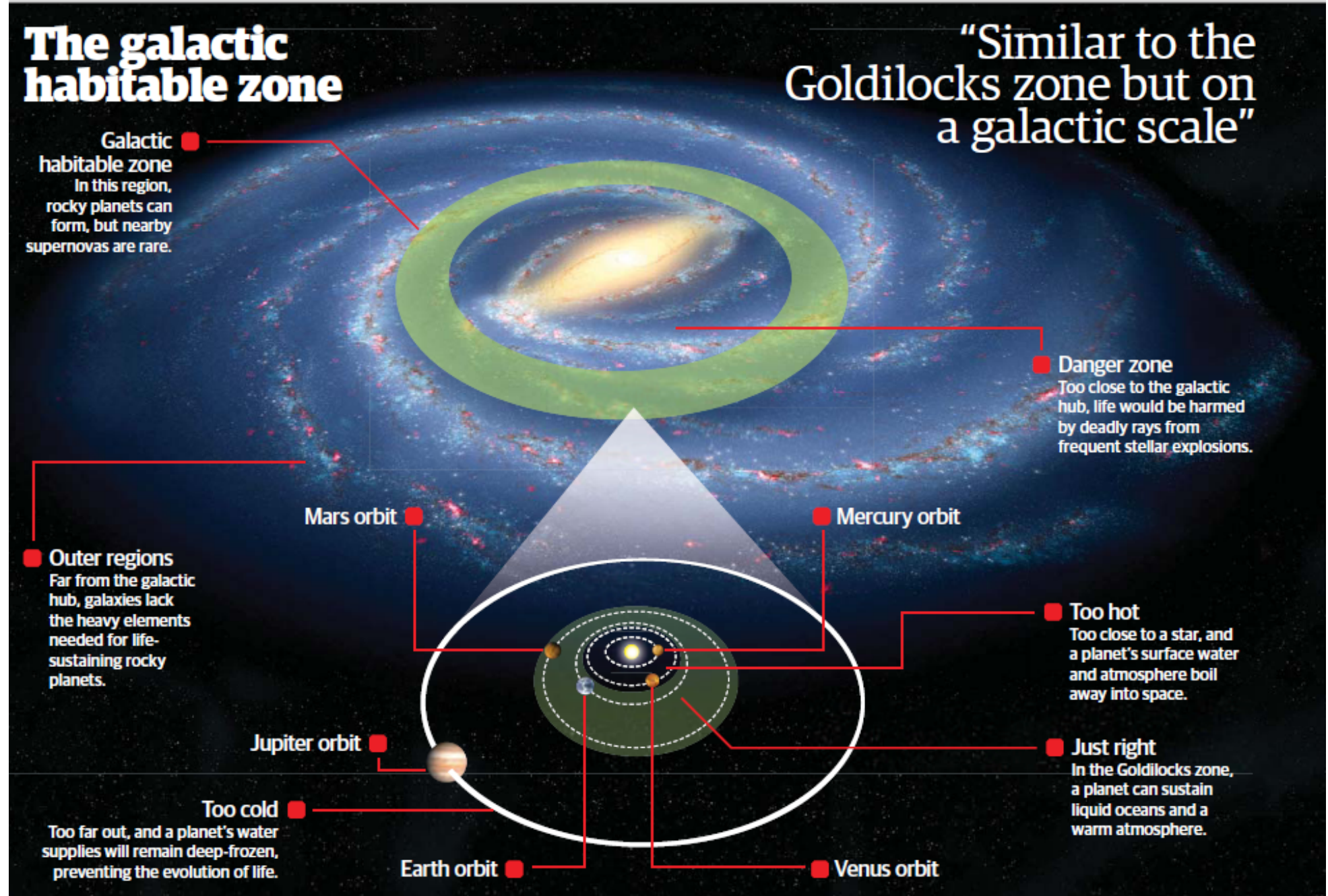
Jupiter orbit

Just right
In the Goldilocks zone, a planet can sustain liquid oceans and a warm atmosphere.

Too cold
Too far out, and a planet's water supplies will remain deep-frozen, preventing the evolution of life.

Earth orbit

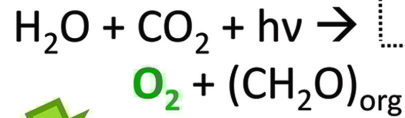
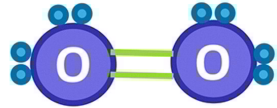
Venus orbit



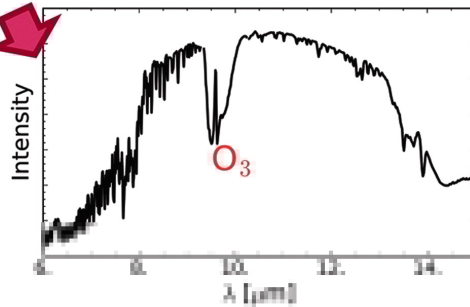
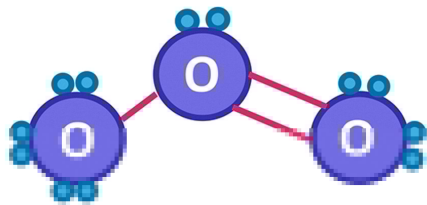
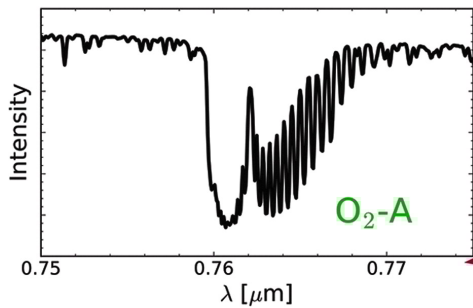
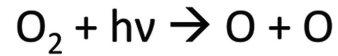
Biosignatures: detecting life elsewhere

Gaseous

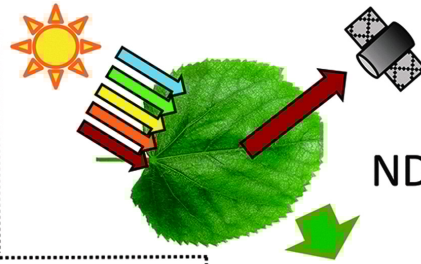
Ex: **Oxygenic Photosynthesis**



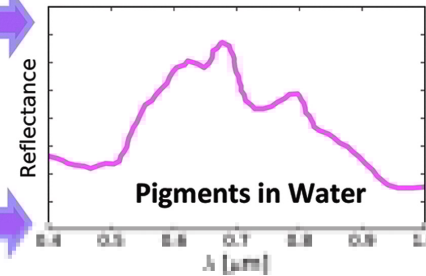
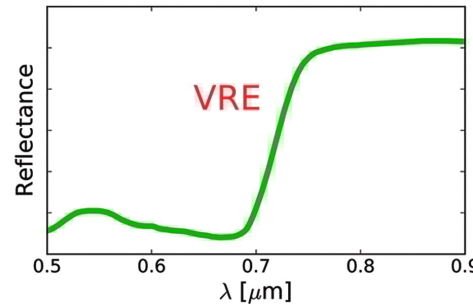
Photolytic Byproduct



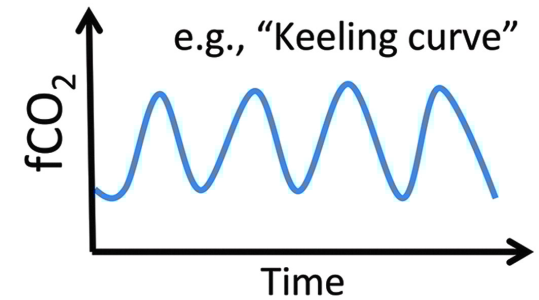
Surface



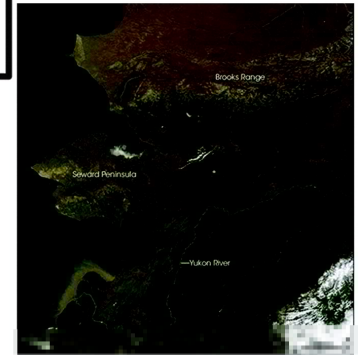
$$\text{NDVI} = \frac{\text{IR} - \text{VIS}}{\text{IR} + \text{VIS}}$$



Temporal



Seasonal Changes in Gases or Surface

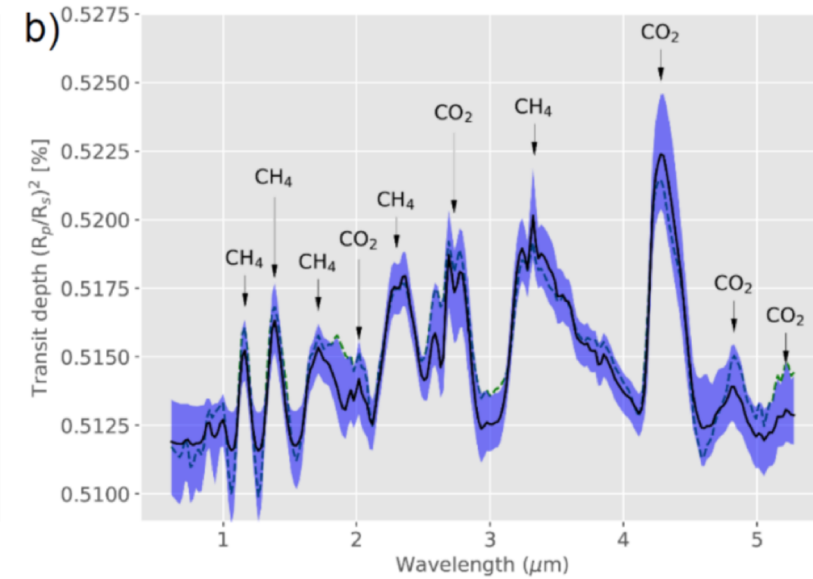
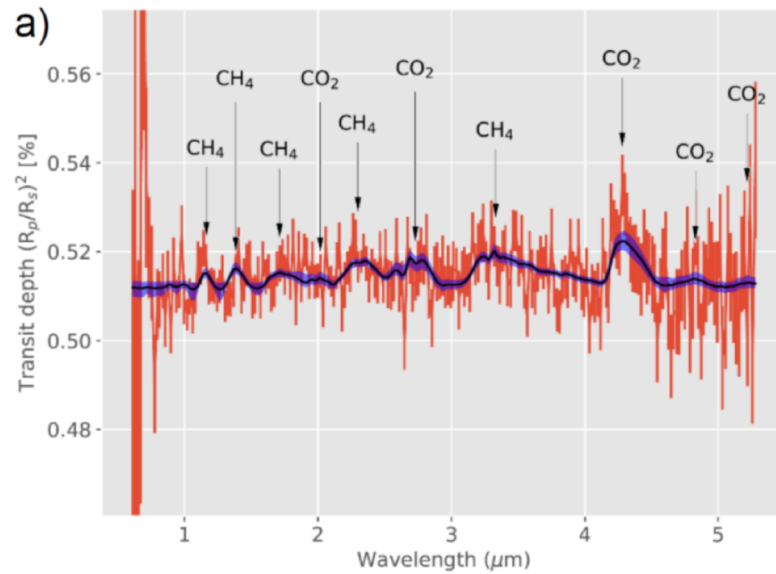


Best prospect for exoplanets: strong chemical disequilibrium detected by spectroscopy

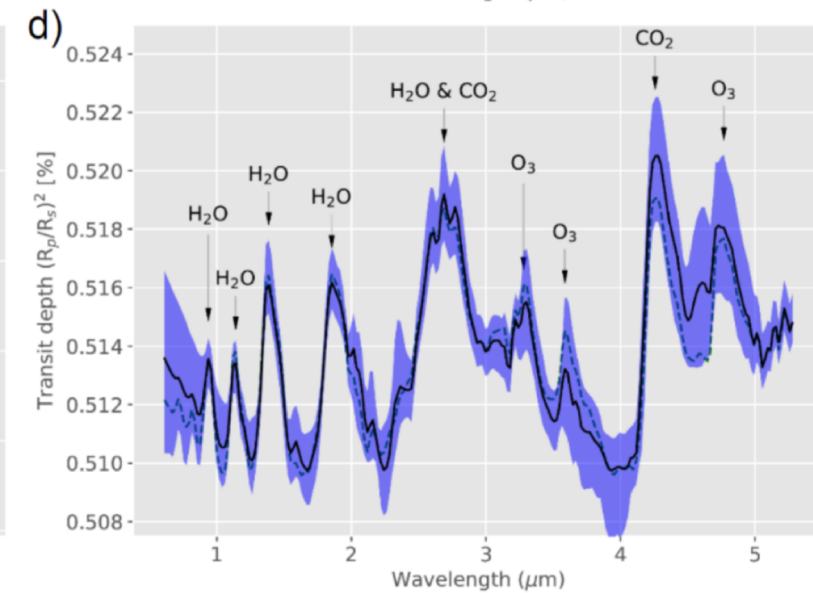
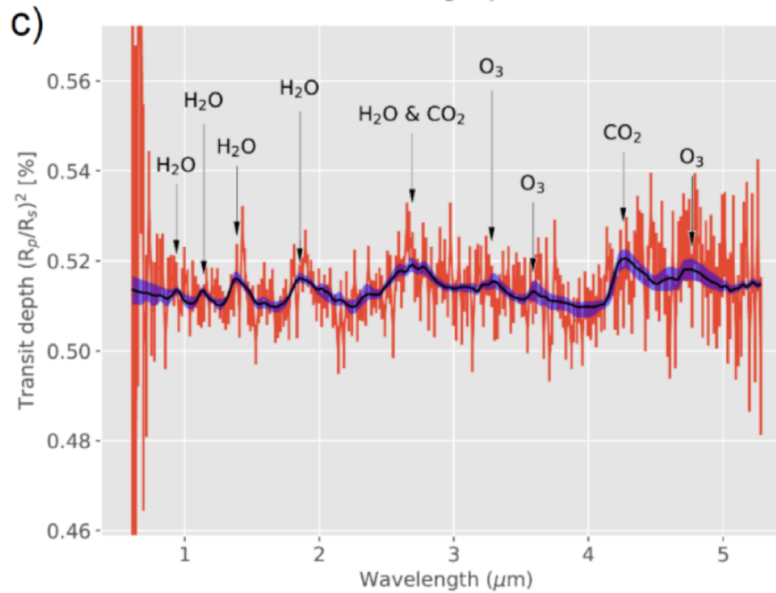
Best shot: O₂ / CH₄. Together, these gazes form quickly CO₂ and H₂O

Case study: TRAPPIST-1e in transmission

Archean
Earth



Modern
Earth



No life around M dwarfs?

Tidal locking

The planets always shows the same face to the star.

→ The atmosphere freezes gradually on the night side (atmospheric collapse)(Haberle et al. 1996)

→ **3D climate modeling show that 0.1bar is enough to avoid this collapse**
(Joshi et al. 1997)

No flux in the visible

So no photosynthesis possible?

Chlorophyll f absorbs light at 706nm (Chen et al. 2010)

Some bacteriochlorophylls absorbs up to 1.2 microns

M-dwarfs are magnetically actives (flares + spots)

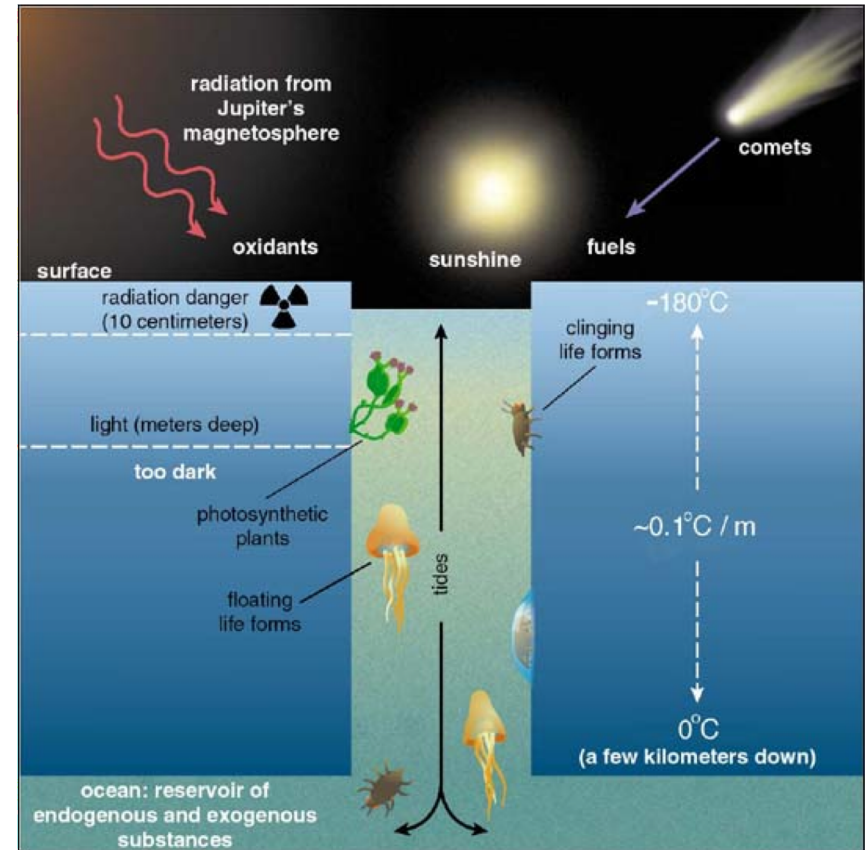
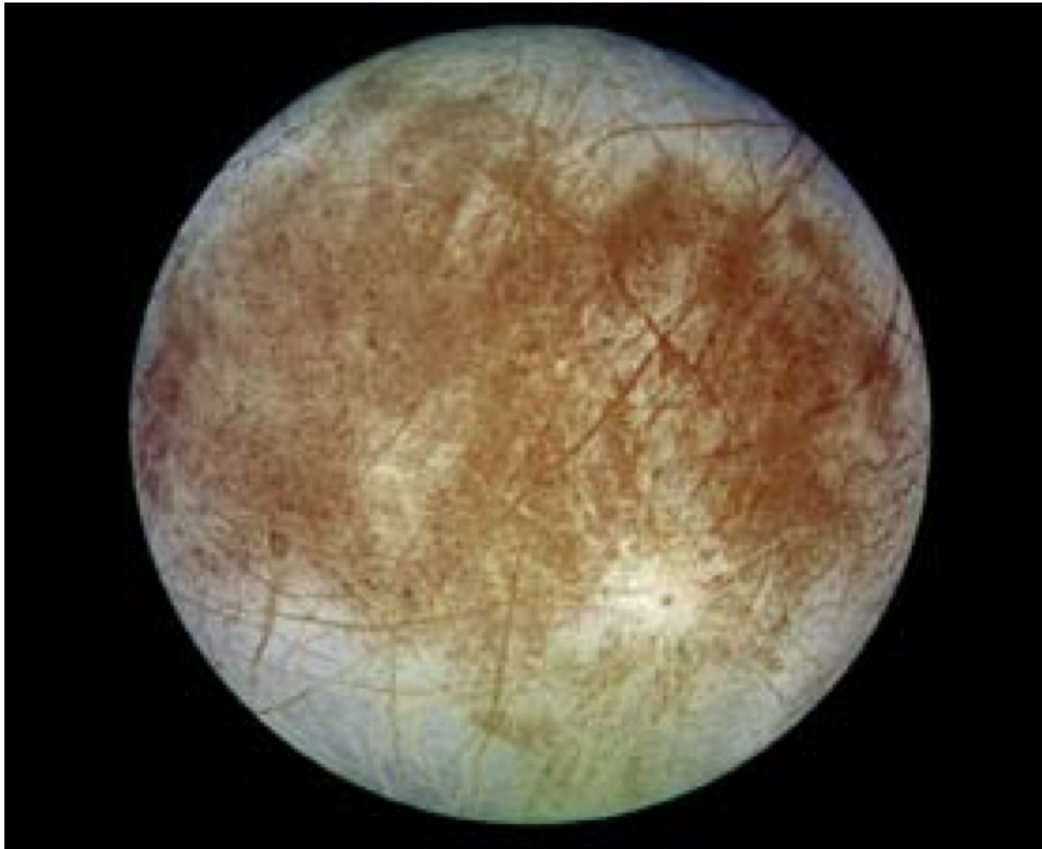
Spots → **effects similar to seasons on Earth** (Joshi et al. 1997)

Flares → **emission in UV and X very strong but sporadic. The cumulated impact on the ground could be similar in amplitude to Earth** (Heath et al. 1999)

→ strong emission of charged particules, so erosion of the atmosphere and harsh biological damages. A magnetosphere is required, but tidal locking could inhibit it (slow rotation).

But Mercury has a 59 days rotation and it has a significant magnetosphere...

No life outside the HZ?



Richard Greenberg

Why not? But for exoplanets, we are forced to focus on biospheres giving rise to strong atmospheric signatures (biosignatures) -> **we focus on HZ**

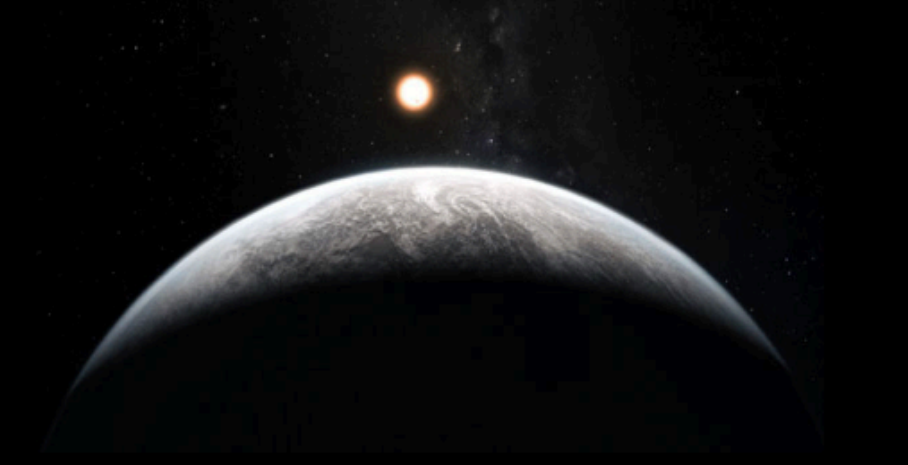
How many potentially habitable planets in the Galaxy?

Dressing & Charbonneau (2013)

Petigura et al. (2013)



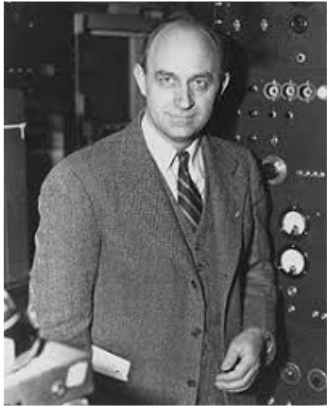
Red dwarfs: 1/6 to 1/7
20 billions



Solar-type stars: 1/5
8 billions

At least 25 billions of potentially habitable planets!
Next step: study the atmospheres of some of the nearest

Exoplanets and Fermi's paradox



Enrico Fermi
(1901-1954)

Summer 1950, Los Alamos: « Where is everybody? »

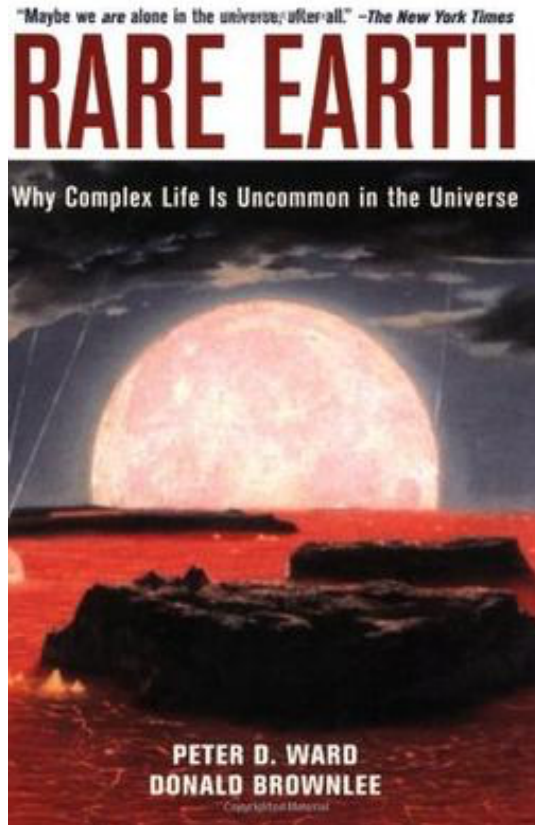
No trace of alien technology in any of our observations of the Universe

2019: Milky way hosts ~300 billions stars, and billions of potentially habitable planets. The Universe contains hundreds of billions of galaxies.

Several classes of solution (Cirkovic 2018):

- *Logistic*: They have not been able to come yet (interstellar travel is really hard)
- *Catastrophist*: They have been destroyed (SN, gamma-ray burst, self-destruction...)
- *Rare Earth*: they do not exist (early great filter, consciousness is ephemeral)
- *Solypсист*: they are here, or we live in a simulation
- *The Great Old Ones*: they have been there for long and do not like the competition

The Rare Earth – Great Filter Hypotheses



Peter D. Ward – paleontologist
Donal Brownlee – astrophysicist

Habitable planets are frequent in the Universe
Chemical elements required by life are frequent in the Universe
Microbial life is probably frequent

BUT

The apparition of multicellular complex organisms –and especially technologically advanced species- requires so many special conditions and unlikely events that the odds for our galaxy to host another civilization are close to zero

Only one low-probability event (**Great Filter**) in the chain of events that led to the apparition of life and its evolution to complex organisms and humans could explain the Fermi's paradox.

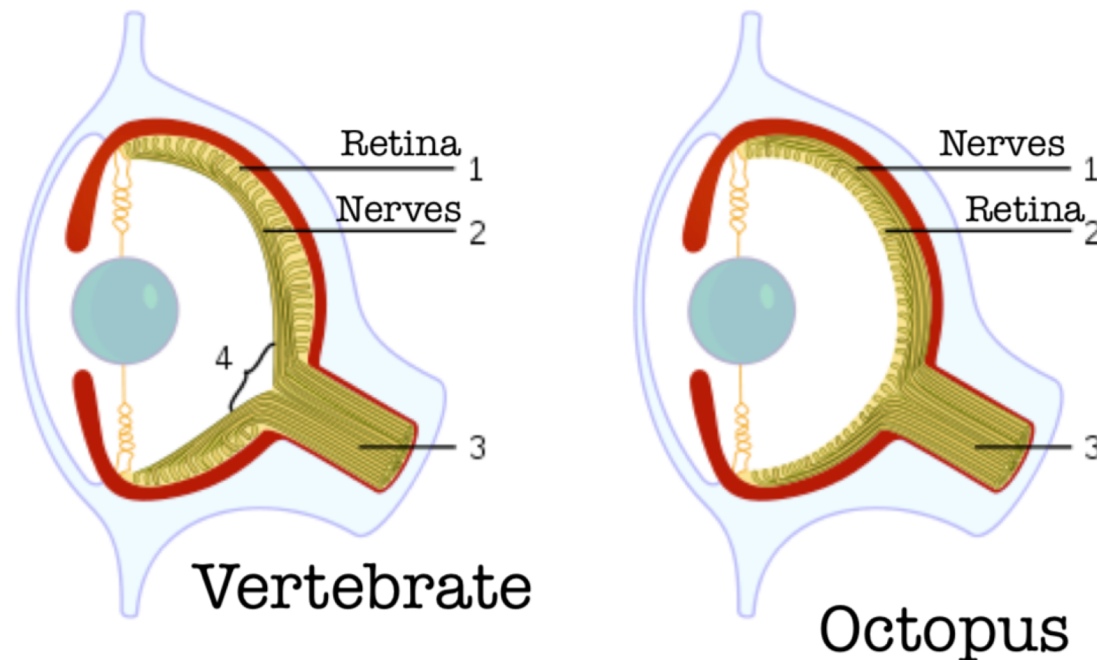
Critics against the Rare Earth – Great Filter Hypotheses

Misconception: « evolution building an intelligent being from microbes is as unlikely as a tornado in a junkyard assembling a 747 ».

BUT: evolution is not a random process, nor is it a process with a « goal » . It is based on the selection of traits produced by mutations that enable a species to survive under the pressure of the environment, including the competition from other species.

-> **global tendency towards complexity.**

Convergence: there are only a limited number of evolutionary solutions that work

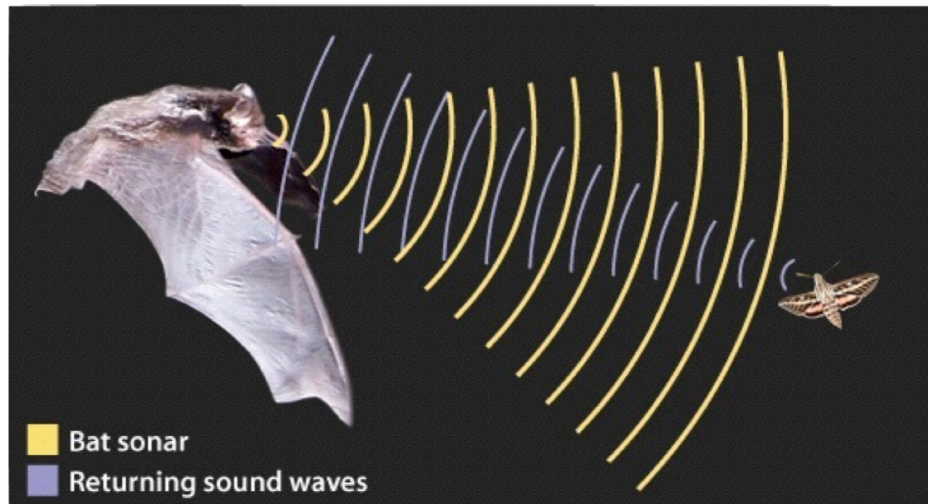


Critics against the Rare Earth – Great Filter Hypotheses

Intelligence is clearly a solution that works (for a time, at least...)



The diversity of the evolutionary « solutions » is much larger than generally thought



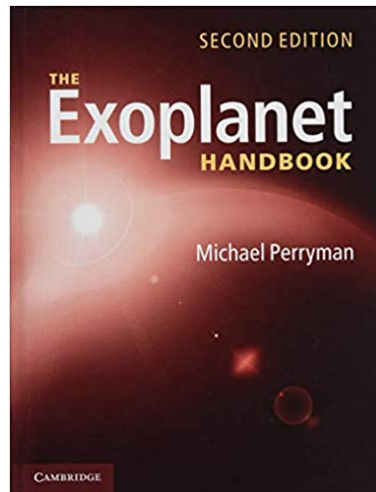
Mantis Shrim: 16 kinds of photoreceptors from UV to deep-red and polarized light

The Great Old Ones

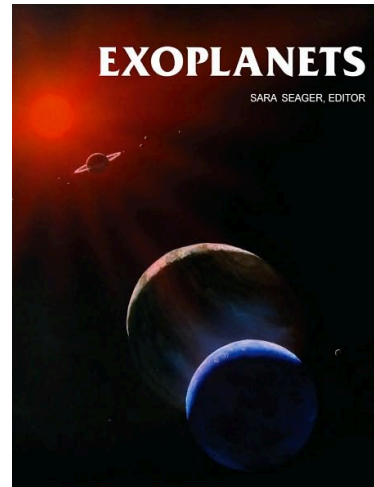


<http://shop.alexanderjansson.com/product/the-great-old-ones>

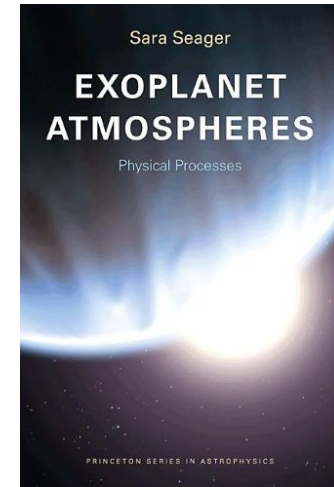
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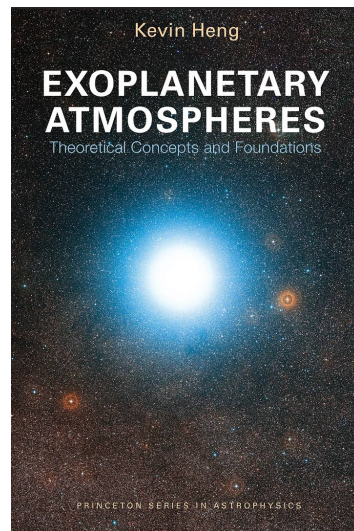
M. Perryman
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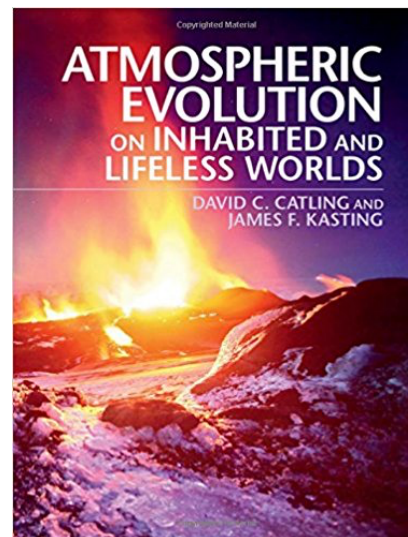
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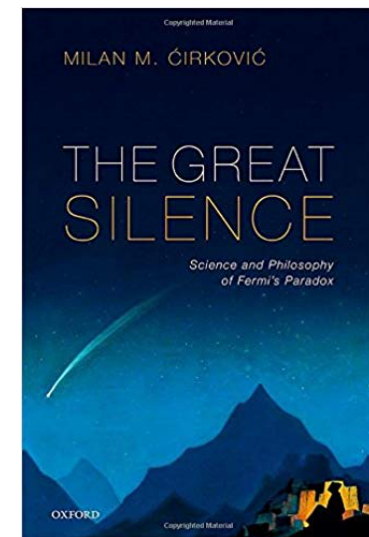
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M. M. Ćirković
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